

Short-term thermal stratification and partial overturning events in a warm polymictic reservoir: effects on distribution of phytoplankton community

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Abstract

In lentic freshwater ecosystems, patterns of thermal stratification play a considerable part in determining the population dynamics of phytoplankton. In this study we investigated how these thermal patterns and the associated hydrodynamic processes affect the vertical distribution of phytoplankton during two consecutive diel cycles in a warm polymictic urban reservoir in the metropolitan region of São Paulo, Brazil. Water samples were taken and physical, chemical and biological data collected at half-meter intervals of depth along a water column at a fixed site, every 3 hours throughout the 48-hour period. Two events of stratification, followed by deepening of the thermocline occurred during the study period and led to changes in the vertical distribution of phytoplankton populations. *Aphanocapsa delicatissima* Nägeli was the single dominant species throughout the 48-hour period. In the second diel cycle, the density gradient induced by temperature differences avoided the sedimentation of *Mougeotia* sp. C. Agardh to the deepest layers. On the other hand, *Pseudanabaena galeata* Böcher remained in the 4.0-5.5 m deep layer. The thermal structure of the water was directly affected by two meteorological factors: air temperature and wind speed. Changes in the cell density and vertical distribution of the phytoplankton were controlled by the thermal and hydrodynamic events.

Keywords: phytoplankton, thermohydrodynamics, thermal stratification, diel variation, Guarapiranga reservoir.

Estratificação térmica de curto prazo e eventos de desestratificação parcial em um reservatório polimítico quente: efeitos sobre a distribuição da comunidade fitoplanctônica

Resumo

Em ecossistemas de água doce, os padrões de estratificação térmica têm uma influência considerável sobre a dinâmica populacional do fitoplâncton. Neste estudo nós investigamos como os padrões de estratificação térmica e os processos hidrodinâmicos afetaram a distribuição vertical do fitoplâncton durante dois ciclos nictemerais consecutivos em um reservatório urbano na cidade de São Paulo. As amostragens e medidas das variáveis físicas e químicas foram realizadas a cada três horas durante um período de 48 horas, ao longo de um perfil vertical. Ocorreram dois eventos de estratificação e subsequente abaixamento da termoclina acarretando alterações na distribuição vertical do fitoplâncton. *Aphanocapsa delicatissima* Nägeli foi a principal espécie dominante durante todo o período avaliado. No segundo ciclo nictemeral o gradiente de densidade criado pelas diferenças de temperatura impediu o afundamento de *Mougeotia* sp. C. Agardh para as camadas mais profundas. Por outro lado *Pseudoanabaena galeata* Böcher se manteve nas camadas de 4.0 e 5.5 m. A estrutura térmica da água foi diretamente afetada pelos fatores meteorológicos temperatura do ar e velocidade do vento. As mudanças na densidade e na distribuição vertical do fitoplâncton foram controladas pelos eventos térmicos e hidrodinâmicos.

Palavras-chave: fitoplâncton, termohidrodinâmica, estratificação térmica, variação nictemeral, reservatório de Guarapiranga.

1. Introduction

Thermal stratification is an important natural phenomenon in aquatic systems, which interferes significantly with their physical and chemical structure, creating complex gradients or simply leading to increased heterogeneity of the water column. The stratification pattern and the temperature of the upper mixed layer depend greatly on the latitude; thus the changes in stability are much more variable in the short term in tropical than in temperate lakes, particularly if winds are strong enough to cause vigorous motion in the upper water layer (Lewis, 1987; Cantin et al., 2011).

Recent research has drawn attention to the fact that the ecological study of plankton communities in lakes and reservoirs requires more than the traditional limnological approach, based on the chemical analysis of the water body and the biology of the component microorganisms. Physical data are also needed, particularly with regard to the thermal hydrodynamics of the water body (MacIntyre et al., 2002), since the stratification pattern is a distinctive feature and a key to understand the dynamics of plankton communities in tropical lakes (Sarmiento, 2012). In other words, those physical processes are the primary causes of first-order effects that then propagate through numerous and more complex effects of an ecological nature (Lewis, 1987).

The thermal hydrodynamics of warm polymictic reservoirs is normally governed by density gradients. According to Reynolds (1984), changes in water density resulting from a rise of 1°C from 24°C to 25°C is about 3 times as large as that observed between 4°C and 5°C, increasing the likelihood of distinct, stable layers being formed in warm regions. However, the heat distribution and thus the density gradients in tropical lakes are constantly changing, these water-bodies being more susceptible to numerous partial mixing events, usually tending to become stratified in the daytime and mixed at night (Lewis, 1987; Sarmiento, 2012).

In temperate climates, the thermal stratification patterns in lakes are mainly associated with seasons (Hutchinson, 1957; Imberger and Patterson, 1989), showing a well-defined sequence of events. By contrast, in tropical regions, where seasons are less marked, lake stratification tends to be associated with a sequence of short-term climatic events (such as intense solar radiation, torrential rain, cold fronts or strong winds), which promote multiple cycles of stratification and partial overturning over short time intervals (Talling, 1966; Barbosa and Padisák, 2002; Becker et al., 2009).

Among the various freshwater communities, phytoplankton is particularly affected by the turbulent mixing processes in the water, and stratification and overturning events have a great importance in several features of algal ecology (Reynolds, 1984; Lopes et al., 2005). The composition, vertical distribution and primary productivity of algal populations are directly affected by the interaction of various factors that modulate the thermal and chemical structure of the water column, notably the

heat flux, wind velocity, density, viscosity and nutrient availability (Cantin et al., 2011).

The conditions in which phytoplankton live are evidently heterogeneous, the most variable environmental factors being temperature, irradiation, water flow and nutrients availability. According to Reynolds (1984), an adequate understanding of phytoplankton ecology depends on the knowledge of habitat heterogeneity, of the phenomena that produce that heterogeneity, and patterns of distribution of the organisms in space and time.

The objectives of this study were to determine how the physical structure of the water column influences the vertical distribution of phytoplankton in a warm polymictic tropical reservoir used for the public water supply. The main questions posed in this study were: i) How do meteorological factors alter the thermal regime in the Guarapiranga reservoir in the short term? ii) How does the phytoplankton community respond to thermal and hydrodynamic events, in terms of changes in population densities and vertical distribution?

2. Material and Methods

2.1. Study area

The Guarapiranga (23°46'S; 46°43'W) is a large subtropical polymictic reservoir (Gemelgo et al., 2009) located in the Upper Tietê River basin (Figure 1). This reservoir has 33 km² surface area, 5.7 m mean depth and 185 d water retention time (Beyruth, 2000). It forms part of the water supply system for the Metropolitan Region of São Paulo, contributing around 20% of the water consumed, and is the second longest freshwater body in this region. Since the 1970s, the local population density has risen and conservation areas around the reservoir have been occupied. This process, coupled with the absence of an effective municipal sewage system in the catchment area of the reservoir, has led to an accelerated eutrophication and a fall in the water quality (Beyruth, 2000).

2.2. Meteorological variables

Air temperature and wind speed and direction were recorded throughout the study period, at intervals of 10 s, by means of a floating weather station located near the dam. Rainfall data were collected every 10 minutes from a public weather station located a distance of 5 km from the sampling point.

2.3. Physical and chemical water variables

Samples of water were taken for chemical analysis from depths of 0.0, 2.0, 4.0 and 5.5 m, near the point of entry of the Embu-Mirim tributary stream into the Guarapiranga reservoir (23°42'53''S and 46°44'32''W), every 3 hours for a 48 h period, from September 12th to 14th, 2010 (dry season). The site was chosen for its frequent cyanobacterial blooms. The samples were frozen and analyzed for total nitrogen and phosphorus contents in the laboratory by ion chromatography according to the methods described by Colina and Gardiner (1999). Temperature, dissolved oxygen concentration, pH and

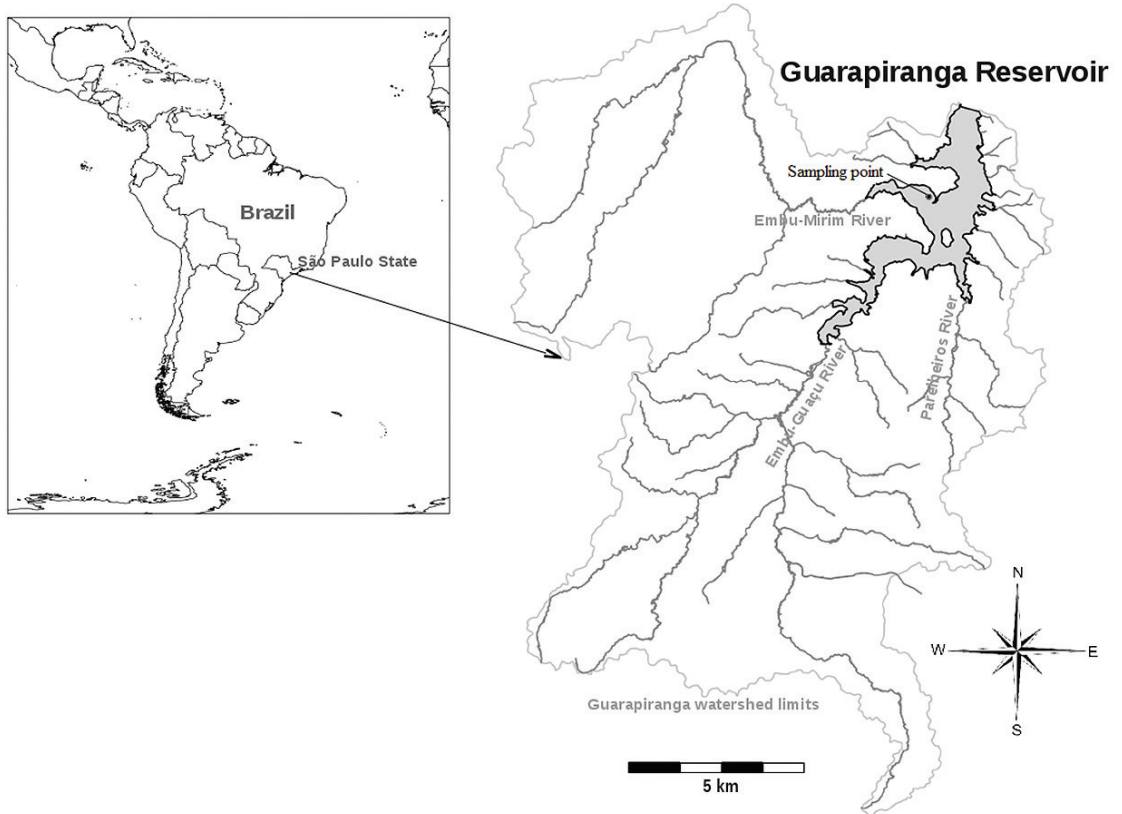


Figure 1. Map of Brazil showing location of Guarapiranga Reservoir, main tributaries and sampling site.

electrical conductivity of the water were measured *in situ*, at intervals of 0.5 m from the surface down to a depth of 5.5 m, with an U-10 Horiba portable multiprobe. The transparency of the water and euphotic zone depth (defined as 1% of surface irradiance) were determined as described by Margalef (1983), from Secchi disk readings.

2.4. Biological variables

A total of 64 phytoplankton samples were collected in 5 L Van Dorn bottles, from depths of 0.0, 2.0, 4.0 and 5.5 m, and fixed with neutral lugol solution. Individual algae were counted by examination in a Zeiss Axiovert inverted microscope, after prior sedimentation of a 10 mL sub-sample in an Utermöhl chamber. The time allowed for sedimentation was three hours per cm depth of the sample in the chamber (Utermöhl, 1958). Counting was carried out in horizontal and vertical transects of the field view and the density was calculated by the formula proposed by Ros (1979) and expressed in number of cells per milliliter (cells mL⁻¹).

Chlorophyll *a* concentrations were determined by spectrophotometric methods, as described in Nush (1980), by hot extraction in 90% ethanol.

2.5. Trophic State Index (TSI)

The trophic state index (TSI) was calculated according to Carlson (1977), modified by Toledo et al. (1983),

considering the total phosphorous concentration values. The following criteria were used to classify the trophic state: oligotrophic – TSI ≤ 44; mesotrophic- 44 < TSI ≤ 54; and eutrophic – TSI > 54.

2.6. Statistical analysis

Analysis of variance (ANOVA post-hoc *t*-test significance accepted when $p \leq 0.05$) was used to test for significant differences in phytoplankton density, between distinct depths, employing the free program PAST 1.79 (Hammer et al., 2001). Principal components analysis (PCA), available in the computer program CANOCO 3.12 (Ter Braak and Šmilauer, 2002) was carried out to summarize the main patterns of variation within the physical and chemical variables in the water and to order the samples with respect to their chemical and limnological characteristics. The data for the dominant phytoplankton species were transformed by the function $\log(x+1)$ to normalize them and then subjected to redundancy analysis (RDA), again employing CANOCO 3.12. The Monte Carlo test, with 4999 random permutations was used to test for significant associations between independent significantly correlated abiotic (water temperature and depth) and biological (chlorophyll *a* concentration and phytoplankton density) variables.

3. Results

3.1. Diel variations of meteorological and water column variables

The fluctuations in wind speed, wind direction and air temperature recorded at the sampling times, are presented in Table 1. At the beginning of the study, at 10:00 am on September 12th, 2010, the air temperature was 20.0 °C and a wind of 1.5 m s⁻¹ was blowing from the South. From 7:00 pm until 7:00 am on September 13th, meteorological conditions changed. The wind speed increased to 4.0 m s⁻¹ and changed direction, now blowing from the North (Table 1). Air temperature also varied widely, the amplitude of variation reaching 14 °C. Higher temperatures of around 30.0 °C occurred on the afternoon of September 13th. Simultaneously, the wind speed varied from 0.0 m s⁻¹ to a strong breeze (4.0 to 5.0 m s⁻¹). No rain fell in this period or in the previous 48 hours.

In the second diel cycle, beginning at 10:00 am on September 13th, the air temperature rose gradually from 17 °C to the maximum recorded value of 29.8 °C at 4:00 pm, then fell steadily until 10:00 pm and warmed up again throughout the night and early morning.

3.2. Physical and chemical characteristics of water column

Water physical and chemical conditions also varied with depth, as shown in Table 2. The pH ranged from slightly basic at the surface to slightly acidic near the bottom (7.15 ± 0.65). Total nitrogen and total phosphorus concentrations were high at all depths ($1,932.84 \pm 264.42 \mu\text{g L}^{-1}$ and $40.8 \pm 11.19 \mu\text{g L}^{-1}$ respectively) and chlorophyll *a* concentrations were also high ($27.55 \pm 8.5 \mu\text{g L}^{-1}$) with lowest values at 5.5 m.

The euphotic zone reached a maximum depth of 3.18 m on September 13th at 4:00 pm, coinciding with

Table 1. Wind speed and air temperature recorded by a floating weather station during the study period, from 10 am on September 12 to 10 am on September 14, 2010.

Day	Hour	Wind speed (m s ⁻¹)	Wind direction (°)	Air temperature (°C)
Sept 12	10:00 am	1.50	180 ↑	20.25
	1:00 pm	1.50	180 ↑	26.45
	4:00 pm	3.67	167 ↑	23.20
	7:00 pm	2.65	180 ↑	20.10
	10:00 pm	0.52	154 ↖	18.40
Sept 13	1:00 am	0		17.20
	4:00 am	0.23		16.60
	7:00 am	0		17.00
	10:00 am	3.93	30 ↙	24.14
	1:00 pm	3.46	54 ↙	28.16
	4:00 pm	2.36	26 ↙	29.85
	7:00 pm	1.24	232 ↗	23.64
	10:00 pm	0		19.09
Sept 14	1:00 am	0		19.55
	4:00 am	4.32	332 ↘	22.40
	7:00 am	2.14	360 ↓	21.50
	10:00 am	4.14	5 ↓	25.12

Wind direction defined clockwise, 0° being wind from North.

Table 2. Mean values of physical and chemical variables and calculated trophic state index (TSI) for Guarapiranga Reservoir, between 10 am, September 12 and 10 am, September 14, 2010.

	0.0 m	2.0 m	4.0 m	5.5 m
Secchi (m)	1.0 ± 0.1	--	--	--
Euphotic Zone (m)	2.8 ± 0.2	--	--	--
Water Temperature (°C)	20.4 ± 0.6	19.7 ± 0.5	18.7 ± 0.5	18.2 ± 0.2
pH	7.8 ± 0.6	7.4 ± 0.4	6.9 ± 0.3	6.3 ± 0.3
Conductivity (μS cm ⁻¹)	102.9 ± 5.2	104.2 ± 5.0	103.4 ± 3.1	101.4 ± 6.1
Dissolved oxygen (mg L ⁻¹)	9.6 ± 1.2	7.6 ± 1.5	3.3 ± 2.5	1.1 ± 1.1
Total nitrogen (μg L ⁻¹)	1912.4 ± 198.8	2076.3 ± 163.9	1883.9 ± 335.7	1854.9 ± 287.1
Total phosphorus (μg L ⁻¹)	42.5 ± 8.7	44.3 ± 10.2	38.3 ± 13.5	38.1 ± 11.4
Chlorophyll <i>a</i> (μg L ⁻¹)	32.4 ± 5.1	34.2 ± 6.2	22.8 ± 6.3	20.8 ± 6.5
TSI	> 62			

the higher surface dissolved oxygen (DO) concentration and warmer water (Figure 2). In the deepest layer (5.5 m below surface), low DO concentrations were recorded, with complete anoxia beyond 5.5 m (Table 2), while at the surface the DO concentration was generally high (e.g. 12.5 mg L⁻¹ at 4:00 pm on September 13th).

The trophic state index (TSI, Table 2) indicated that Guarapiranga Reservoir could be classed as eutrophic during the study period.

3.3. Diel changes in the thermal, chemical and biological structure of the water column

Changes in the thermal, chemical and biological profiles of the water column were analyzed in detail at a fixed site over two consecutive diel cycles starting at 10 am on September 12th, 2010.

The water column was initially characterized by moderate stratification, with a temperature gradient of less than 1°C in the 1.0 m thick epilimnion and a moderately

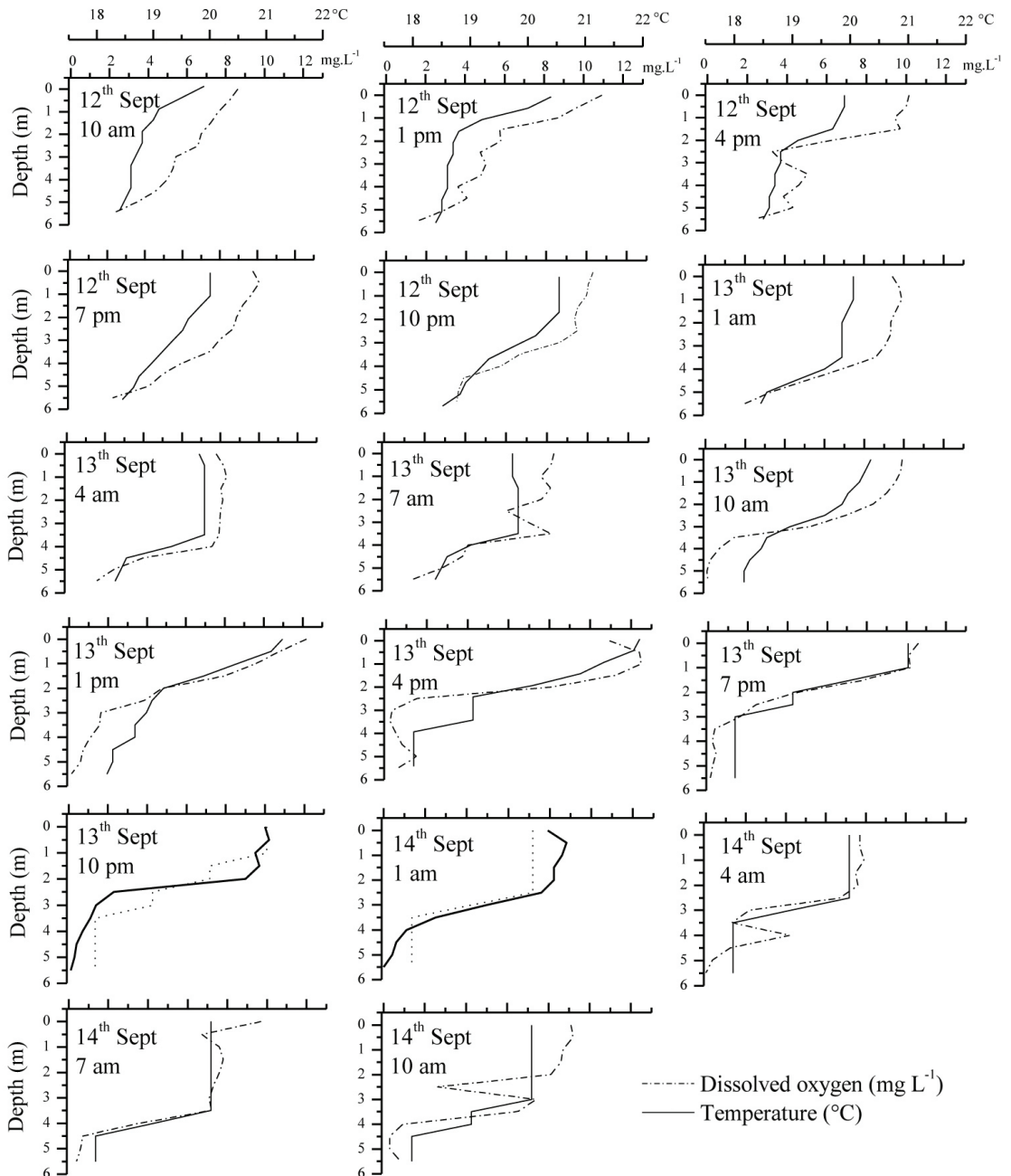


Figure 2. Profiles of water temperature (°C) and dissolved oxygen concentration (mg L⁻¹) against depth at a fixed site in Guarapiranga Reservoir (SP, Brazil), at intervals of 3 h from 10 am on the 12th to 10 am on the 14th of September, 2010.

well defined metalimnion. During the day it progressively formed a stepped thermal structure, returning after 4:00 pm to moderate stratification.

During the night time, with more or less stable temperatures and little wind, the stepped structure was eroded, while the thermocline deepened to 3.5 m increasing the thickness of the mixing layer. A homogeneous and deep epilimnion (down to 3.5 m) appeared in the profiles from 1:00 am till 7:00 am on September 13th.

A third period between 10:00 am and 10:00 pm on September 13th exhibited the same sequence of heat penetrating the upper layers and generating the stepped pattern in the thick metalimnion. From 1:00 pm on September 13th the air temperature rose from 17 °C to 29.8 °C (at 4:00 pm), leading to the strongest thermal stratification recorded in the period of study. At 4:00 pm, the water temperature in the epilimnion was also at its highest (21.8 °C) and there was the steepest gradient (3.7 °C) in the metalimnion (Figure 2). Meanwhile, the depth of the metalimnion at 4.0-5.0 m in the early hours of September 13th, to reach 2.0 m at 1:00 pm, where it stays until 7:00 pm.

After the wind ceased, around 10:00 pm a new thermo-hydrodynamic event started, characterized by thermocline deepening and this extended until the end of the study, at 10:00 pm on September 14th.

3.4. Phytoplankton composition and dominant species vertical distribution

In the period of this study, 68 taxa of algae and cyanobacteria were identified in the phytoplankton community in Guarapiranga, most belonging to Chlorophyceae (41%), Cyanobacteria (25%) and Bacillariophyceae (12%).

In terms of numerical abundance, the dominant species was the cyanobacterium *Aphanocapsa delicatissima* Nägeli. Also abundant were the chlorophyceans *Scenedesmus ecornis* (Ehrenberg) Chodat, *Eudorina elegans* Ehrenberg, *Monoraphidium contortum* (Thuret) Komárková-Legnerová and *Mougeotia* sp. C. Agardh. Species contributing more than 5% of the total phytoplankton density were included in the analysis of changes in the vertical distribution during the 48-hour period. Together, these species represented more than 78% of the total abundance of phytoplankton.

The highest densities occurred at depths 0.0 and 2.0 m and the lowest at 5.5 m (Figure 3). *A. delicatissima* was numerically dominant throughout the study period at almost all sampling depths and time intervals. An exception was the sample taken at 4.0 m depth, on September 13th at 7:00 am, when *Scenedesmus ecornis* was the dominant species. This chlorophycean was the most abundant species among the Chlorophyta in the majority of samples analyzed (Figure 3).

Pseudanabaena galeata was another cyanobacterium occurring in all samples, except on September 13th at 7:00 am. This species occurred in noticeable densities only at 5.5 m, although a few cells were also found in the sample from 4.0 m.

Synechococcus sp. occurred at low densities throughout the study period. With the rising temperatures and strong

winds that characterized the daytime of the second diel period (7:00 am to 7:00 pm September 13th), the *Synechococcus* sp. abundance increased approximately eight times.

The distribution of *Cryptomonas* sp. was heterogeneous in the water column, but the highest density values were always recorded near the surface (0.0 m). A more homogeneous distribution was only observed at 1:00 am on September 13th, 2010.

Similarly to *Cryptomonas* sp., *Eudorina elegans* remained in the upper layers (0.0 and 2.0 m). In the period between 1:00 am and 7:00 am, on both days, *E. elegans* densities increased around four times.

As for most species of algae occurring in Guarapiranga Reservoir, *Mougeotia* sp. was abundant at 0.0 and 2.0 m depths. Only at 7:00 and 10:00 pm on September 12th and 4:00 am on September 13th were higher densities of this species recorded in the deepest layer (5.5 m).

Following the same pattern *Monoraphidium contortum* was also more abundant in the upper layers (0.0 and 2.0 m), except at 4:00 am on September 13th, when like *Mougeotia* sp. it was more numerous at greater depth (Figure 3A)

3.5. Short-term changes in vertical distribution of whole phytoplankton community

Phytoplankton density and chlorophyll *a* concentration differed among water layers (depths) and sampling times following changes in the thermal structure. The vertical distribution of the whole phytoplankton community and temperature profiles are presented together, for each time interval sampled, in Figure 4.

In the first period, from 10:00 am to 10:00 pm on September 12th, algal populations were quite evenly distributed in the water column. Mean densities were 5,917 ($\pm 1,184$) cells mL⁻¹; 6,071 ($\pm 1,674$) cells mL⁻¹; 4,507 ($\pm 1,019$) and 5,965 cells mL⁻¹ ($\pm 1,399$) at 0.0; 2.0, 4.0 and 5.5 m respectively. There were no significant differences in phytoplankton densities between the sampled depths ($F = 1.526$; $df = 3$; $p = 0.2461$).

In the second period, characterized by overturning in the epilimnion and deepening of the thermocline, from 1:00 am until 7:00 am on September 13th phytoplankton distribution was heterogeneous with high densities in the two upper layers, with 10,419 ($\pm 1,541$) cells mL⁻¹ and 8,442 ($\pm 1,431$) cells mL⁻¹ at 0.0 and 2.0 m respectively, and low in the deeper layers, with 5,025 ($\pm 1,429$) and 4,100 ($\pm 2,237$) cells mL⁻¹ at 4.0 e 5.5 m respectively. The ANOVA analysis showed significant differences between the total phytoplankton density at 0.0 m and those recorded at 2.0, 4.0 and 5.5 m ($F = 9.09$, $df = 3$; $p = 0.005$).

The following period (from 10:00 am to 10:00 pm on September 13th) was characterized by strong thermal stratification and a rise and fall in the thermocline. Phytoplankton was concentrated mainly at the surface where the density was 11,790 ($\pm 2,272$) cells mL⁻¹. Densities at other depths were 7,667 ($\pm 1,509$) cells mL⁻¹ at 2.0 m; 5,574 (± 800.6) at 4.0 m; and 5,036 (± 689.0) cells mL⁻¹ at 5.5 m. There were significant differences between depths ($F = 15.91$, $df = 3$; $p = 0.0001$), and the phytoplankton

Short term thermal stratification and phytoplankton

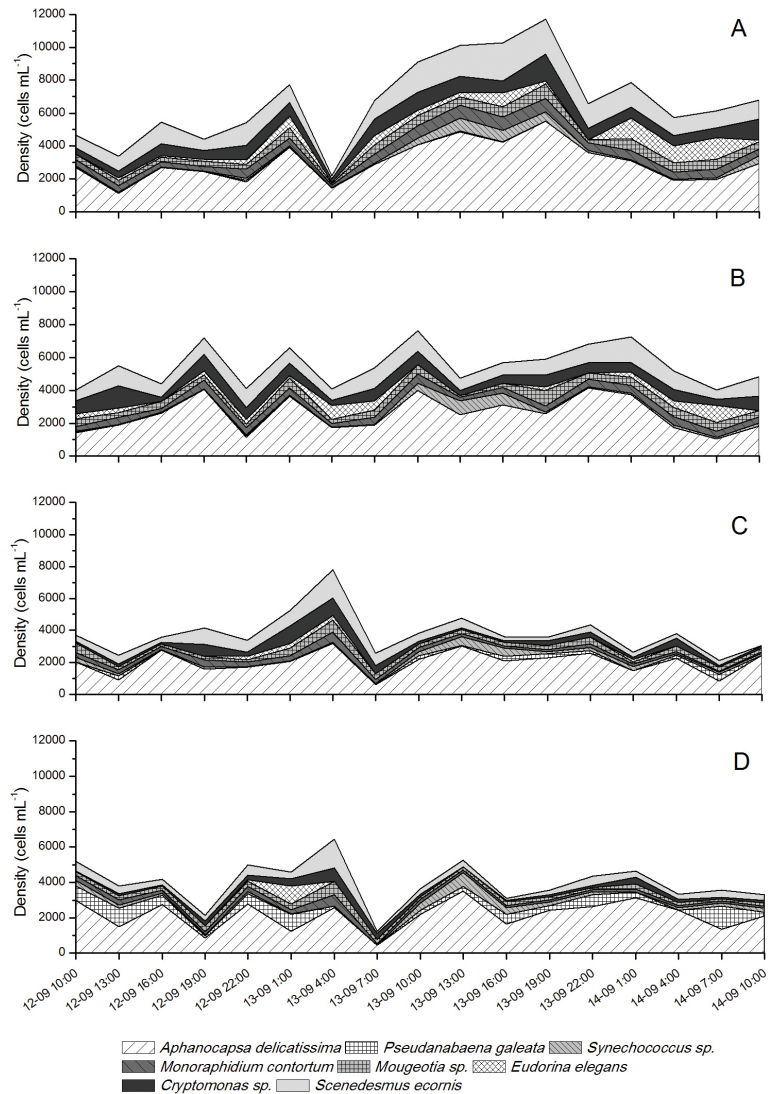


Figure 3. Cell density (cells mL⁻¹) of the dominant phytoplankton species in Guarapiranga Reservoir (SP, Brazil), at various depths (A: 0.0 m, B: 2.0 m, C: 4.0 m and D: 5.5 m), each 3 h from 10:00 am on the 12th to 10 am on the 14th of September, 2010.

density at 0.0 m was statistically different from that at 2.0, 4.0 and 5.5 m depths ($p < 0.05$).

The last period (from 1:00 am to 10:00 am on September 14th), was characterized by night cooling and a deep thermocline, resulting in the redistribution of phytoplankton in the extended epilimnion. The highest mean density was found at 0.0 m ($8,396 \pm 332.3$ cells mL⁻¹) and the lowest at 4.0 m ($4,299 \pm 349.7$ cells mL⁻¹). Statistical tests indicated significant differences between depths ($F = 9.41$, $df = 3$; $p = 0.001$). Phytoplankton densities at 0.0 m were statistically different from those at 4.0 and 5.5 m ($p < 0.05$) and the densities at 2.0 m also differed significantly from that at 4.0 m ($p = 0.046$).

3.6. Ordination analysis

PCA analysis was performed to study the relationships between abiotic variables, water column depth and time

intervals (Figure 5). There were strong relationships, as evidenced by the high explicability attained by the first two axes, which together explained 84.6% of the total variance. The first axis showed negative correlations between all abiotic variables and depth. This axis separated the surface and 2.0 m samples on the left side and 4.0 and 5.5 m depth on the right side. As expected, water temperature, dissolved oxygen concentration and pH were always higher in the surface layers, whereas the total nitrogen, total phosphorus and electrical conductivity, although also predominantly related to the surface layers, were sometimes related to the layer at 4.0 m.

The RDA performed to reveal associations between the population densities of the dominant phytoplankton species and selected abiotic variables is shown in Figure 6. Phytoplankton population densities were strongly associated with the variables water temperature ($p < 0.05$), chlorophyll

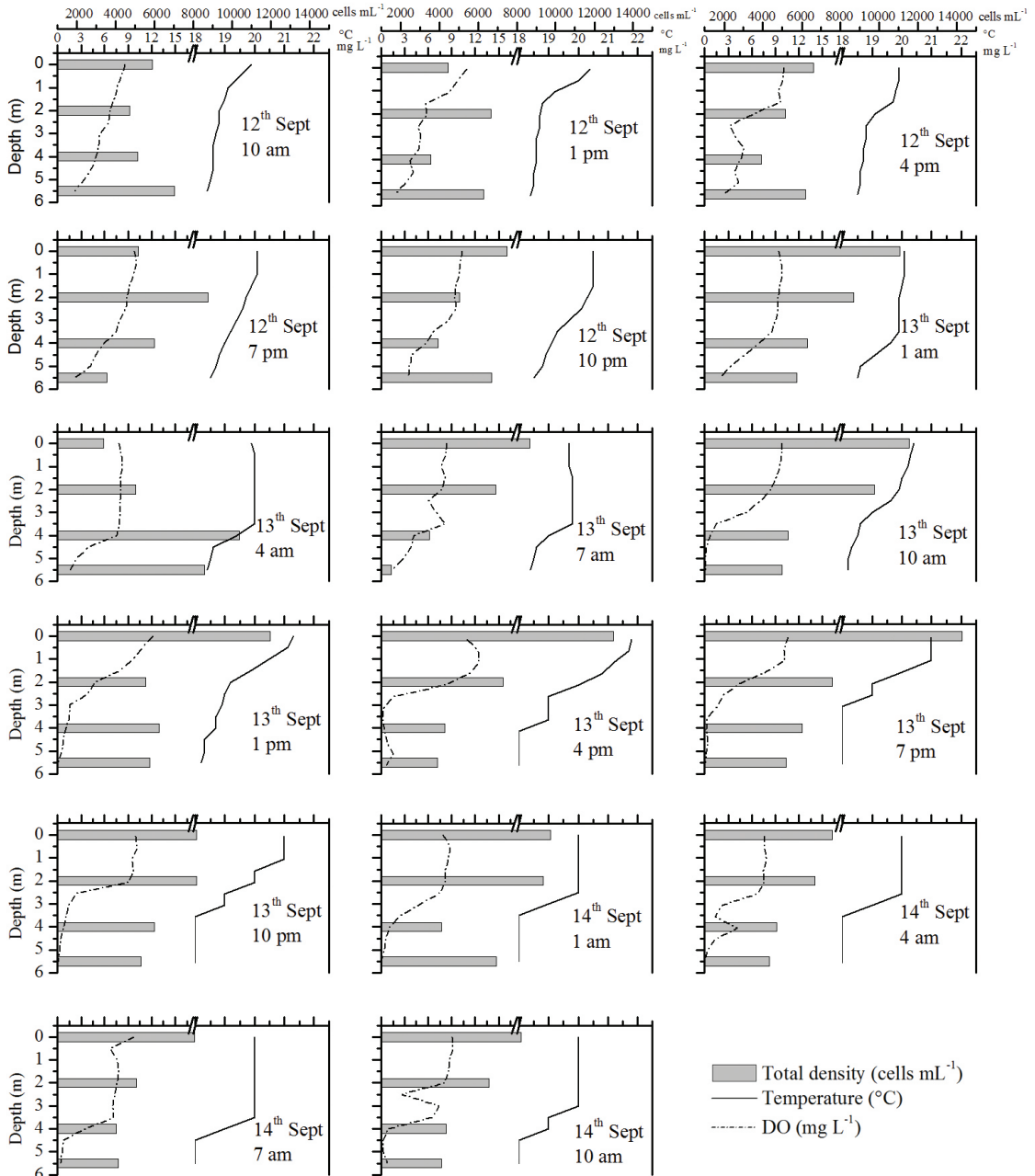


Figure 4. Profiles of water temperature ($^{\circ}\text{C}$), dissolved oxygen concentration (mg L^{-1}) and total phytoplankton density (cells mL^{-1}) measured at various depths in Guarapiranga Reservoir (SP, Brazil), at intervals of 3 h from 10 am on the 12th to 10 am on the 14th of September, 2010.

a ($p < 0.05$) and depth and the first two axes explained 96.3 and 3.4% of the variance, respectively. Population densities of the chlorophytes *Scenedesmus ecornis*, *Monoraphidium contortum*, *Mougeotia* sp. and *Eudorina elegans*, the cryptophyte *Cryptomonas* sp. and the cyanobacterium *Aphanocapsa delicatissima* were positively associated with the depths 0.0 and 2.0 m and also with high water temperature and chlorophyll a concentration. On the other hand, population densities of the cyanobacterium

Synechococcus sp. and of *Pseudanabaena galeata* were negatively related to water temperature and chlorophyll a concentration and positively related to water column depth, since they had high densities in the deep layers at 4.0 and 5.5 m.

4. Discussion

During the period of study, the Guarapiranga Reservoir was subjected to two consecutive cycles of stratification

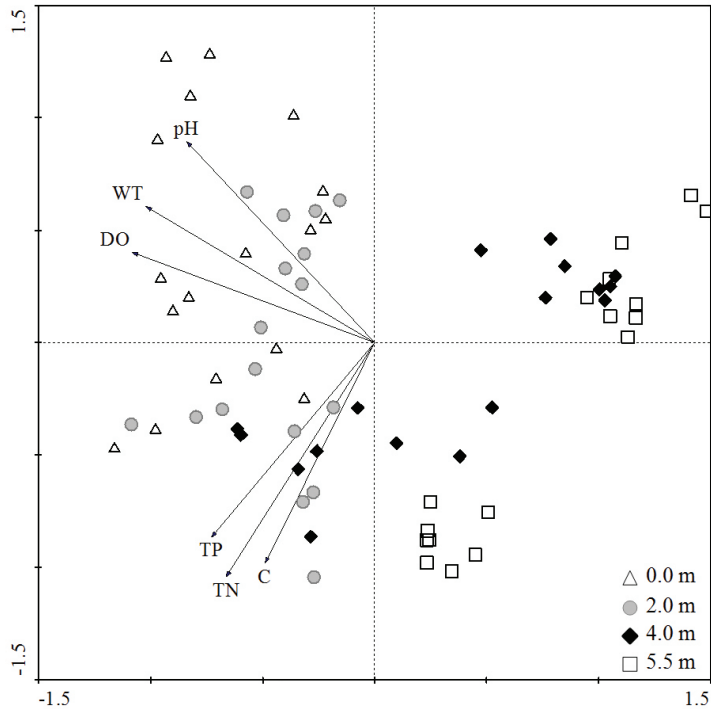


Figure 5. Principal components analysis (PCA) ordination diagram relating the physical and chemical variables with water column depths: WT = water temperature; DO = dissolved oxygen concentration; C = water electrical conductivity; TP = total phosphorus; TN = total nitrogen and pH.

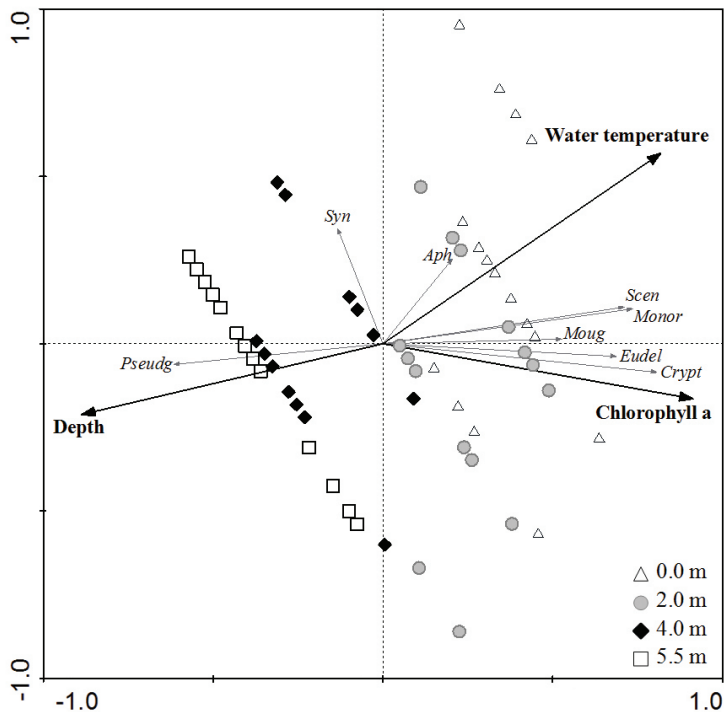


Figure 6. Redundancy analysis (RDA) ordination diagram relating the vertical distribution of dominant phytoplankton species to environmental variables (depth, water temperature and chlorophyll *a*) and to the water column layers. (Aph = *Aphanocapsa delicatissima*; Crypt = *Cryptomonas* sp.; Eudel = *Eudorina elegans*; Monor = *Monoraphidium contortum*; Moug = *Mougeotia* sp.; Pseudg = *Pseudanabaena galeata*; Scen = *Scenedesmus ecornis* and Syn = *Synechococcus* sp.).

followed by partial mixing of the water column. Day-time heating created new temperature gradients in the upper layer resulting in a multi-stepped stable metalimnion. During the night, as surface water cooled, the density increased making this layer heavier than the one beneath it and causing convectional mixing and thermocline deepening.

In the early hours of the day, the surface layer became well mixed, mainly by the penetrative convection associated with the loss of heat from the surface layer during the night. A thermocline could be clearly distinguished at the base of this well-mixed layer. As the air temperature rose during the morning, so did that of the water surface layer. The thermocline remained shallow, at a depth of 2.0 m or less, owing to heat diffusion within the epilimnion. The wind blew throughout the day, causing turbulent kinetic energy in the surface layer, reinforcing the downward transport of heat, lowering the water temperature near the surface and deepening the thermocline. New thermoclines started forming in the deeper water, within a stepped metalimnion of thermoclines accumulated during net heating of the surface layer, a process commonly seen in tropical lakes (Lewis, 1973; Barbosa and Padisak, 2002).

The wind blowing during the day was strong on the first day, generating moderate stratification and a more even distribution of phytoplankton in the water column. During the night, with a stable and deeper thermocline, the highest concentrations of phytoplankton were found in the upper layers.

In Payne's view (1986), the metalimnion may act as a second bottom for the water column, where the sedimentation of particles may be prevented by its marked density gradient. The metalimnion in the Guarapiranga Reservoir water column indeed acted as a second bottom keeping the chlorophytes *Scenedesmus ecornis* and *Eudorina elegans* practically isolated from the filamentous cyanobacterium *Pseudanabaena galeata*, which was only found in the hypolimnion.

The second day was marked by stronger stratification, determined by a greater heat influx and more constant wind speed. Phytoplankton in the upper layers appeared to have favorable conditions for growth during the second diel cycle, since light and temperature were adequate and nutrients were not limiting.

This could be an explanation for the marked increase of *Aphanocapsa delicatissima*, whose density almost doubled over the first 12 hours. According to the literature, a species of this genus, *A. elachista*, has a doubling time of approximately 15 hours (Rippka, 1972).

It may be hypothesized that at least part of the increase in *A. delicatissima* density and that of some co-dominant algae could be due to population growth, implying a biological response of phytoplankton replication in a shorter time than the range of 50-200 hours suggested by Harris (1986).

The high availability of nutrients in Guarapiranga Reservoir was noted in previous work (Gemelgo et al., 2009) in which the inorganic nitrogen compounds and

total phosphorus were recorded, and also from the total nutrients measured in the present study. During that study the temperature and euphotic zone were around 20 °C and 3 m, respectively.

Most species remained in the upper 0.0 and 2.0 m depth. In some cases, this can be related to the small cells size (*Scenedesmus ecornis* and *Synechococcus* sp.) or to the buoyancy of the colony mucilaginous matrix (*Aphanocapsa delicatissima*) that avoid sedimentation (Barbosa et al., 2011). Even large filamentous species such as *Mougeotia* sp. were held in the upper layer for long periods. This alga was possibly detached from floating macrophyte periphyton by the water movements in the upper layer. During the morning and early afternoon of the second diel cycle, air temperature increased by 5.6 °C and net heat flux was high, with a 1.4 °C rise in the upper layer. *Mougeotia* sp. filaments were kept in the well mixed epilimnion by the strong density gradient. This could also be the case of the large flagellated *Eudorina elegans* and *Cryptomonas* sp. that remained mainly at the surface. Although these species are able to swim against gravity, according to Reynolds et al. (1983) the fastest velocities they can attain (usually less than 1 mm s⁻¹) are slower than the turbulent velocities generated by a wind of 1 m s⁻¹. In both diel cycles there were light winds of 3-4 m s⁻¹, which can develop turbulence that far exceeds the speed of most motile algal species.

Pseudanabaena galeata was the only one of the abundant species that appeared only in the deep layers (4.0 and 5.5 m) of Guarapiranga Reservoir. Other studies (e.g. Barbosa et al., 2011) also found this species in the deepest layer of a tropical stratified lake (Lake Carioca, MG, Brazil). It is known that this species preferentially occupies the deeper layers, owing to its tolerance to light deficiency (Reynolds et al., 2002).

In conclusion, air temperature and wind speed were the most important meteorological factors altering the thermal structure of Guarapiranga Reservoir during the two-day period, promoting short-term diurnal patterns of stratification and night-time mixing.

The vertical distribution of the main phytoplankton species was related to the thermal stratification and hydrodynamic events. During periods of moderate stratification and low winds, both small and large cells, colonies or filaments were more evenly distributed in the water column, whereas during periods of thermal stratification all species, except *Pseudanabaena galeata*, were trapped in the upper layers by the strong density gradient and extended metalimnion.

Our findings show that in tropical polymictic reservoirs the hydrodynamic events associated with thermal stratification and water-layer mixing do determine most patterns of vertical distribution of algal populations. Wind speed and direction are important variables that must be included for a better understanding of the factors controlling phytoplankton vertical distribution and consequently their physiological functional performance.

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