

A cathodoluminescence-assisted LA-ICP-MS study of topaz from different geological settings

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ABSTRACT: In this study, the crystallization dynamics and the fingerprints of two topaz crystals from the Ouro Preto region (Brazil), two from the Eastern Brazilian Pegmatite Province (EBPP), one from the Western part of the United States (US) and one from Pakistan were characterized. The combination of scanning electron microscope-cathodoluminescence (SEM-CL) and laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) data allowed for the characterization of these samples. Samples from the Ouro Preto region showed primary crystallization and recrystallization processes, which were demonstrated by a CL-heterogeneous core and CL-homogeneous rims in the SEM-CL images. Their fingerprint was composed of Ca, Ti, V, Cr, Mn, Fe, Cu and Zn. Samples from the EBPP showed both dark and CL-homogeneous SEM-CL images, implying slow and late-stage crystallization, typical of pegmatites. Their fingerprints were composed of Li, Nb, LREE, Ta and W, and of Ti, Nb and HREE. A sample from a Western part of the US displayed CL-heterogeneity with growth and resorption events and well-preserved growth zones, which were noticeable through the SEM-CL images. Its fingerprint was composed of Li, Ti, V, Mn, Nb, LREE, Ta and W. A sample from Pakistan showed an evident metamorphic recrystallization process, which was visible by the incipient and CL-homogeneous luminescence given by the SEM-CL images. Its fingerprint was composed of Ca, Cr, V, Zn, LREE, HREE and W.

KEYWORDS: topaz; trace elements; cathodoluminescence; crystallization dynamics; fingerprint.

INTRODUCTION

Topaz is a F/OH-bearing nesosilicate, $\text{Al}_2\text{SiO}_4(\text{F,OH})_2$, that is usually found as an accessory and late-stage crystallisation mineral in rocks from different geological settings. These include F-rich felsic magmatic rocks (rhyolites and granites), late-magmatic rocks (pegmatites) and post-magmatic rocks (greisens), hydrothermal veins and, finally, in high-temperature and high- to ultrahigh-pressure metamorphic rocks, specifically kyanite topaz-bearing quartzites (Zhang *et al.* 2002, Marshall & Walton 2007, Agangi *et al.* 2014, 2016).

The main chemical composition of topaz, with regard to its major elements, is constant, and its chemical variation is only related to its OH/F concentration ratio (Skvortsova *et al.* 2013). With respect to the trace element concentrations of topaz in other studies, Ti, V, Cr, Mn, Fe, Cu, Zn, Nb, LREE, HREE, Ta and W showed relevant variations

(Leroy *et al.* 2002, Wasim *et al.* 2011, Wu *et al.* 2011, Breiter *et al.* 2013, Agangi *et al.* 2014, 2016, Gauzzi *et al.* 2018).

In order to characterize the crystallization dynamics of the different topaz crystals, two techniques are essential: SEM-CL and LA-ICP-MS. The first one enables researchers to get to know the structural and crystallization features of the studied topaz crystals with images obtained by the SEM-CL. In this case, cathodoluminescence (CL) allows for the visualization of features related to crystal growth, dissolution, replacement, deformation and recrystallization of the different topaz crystals, and, thus, permits the possibility of differing them. The second one allows for the characterization of the chemical signature of the already mentioned crystals with LA-ICP-MS, which will enable a quantitative elemental analysis of them, namely their trace elements. These chemical data will complete the information obtained by SEM-CL images, and will allow for researchers to distinguish between the particularities of the

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crystallization dynamics, as well as the fingerprint of each studied topaz crystal.

GEOLOGICAL SETTING

The topaz samples used in this study were from a wide variety of provenances (Fig. 1): two from the Ouro Preto region (Minas Gerais state, Brazil), specifically Dom Bosco and Antônio Pereira mines; two from the Eastern Brazilian Pegmatite Province, namely from Toca da Onça and Ponto de Marambaia (both located in Minas Gerais state) pegmatite bodies; one from Thomas Range (Utah, the United States); and the last one from Ghundao Hill (Mardan District, Pakistan).

Dom Bosco and Antônio Pereira mines (Ouro Preto, Minas Gerais, Brazil)

The Ouro Preto region is located in the Southeastern part of the state of Minas Gerais, Brazil, and is geologically related to the Quadrilátero Ferrífero District. In this region, the topaz mineralizations occur essentially in the Northeastern area of the Dom Bosco Syncline and in the Northeastern flank of the Mariana Anticline (Dorr II 1969, Gandini 1994, Morteani *et al.* 2002, Rojas 2008, Rojas *et al.* 2009, Gauzzi *et al.* 2018).

The Dom Bosco mine is geologically related to the Northeastern area of the Dom Bosco Syncline hinge zone and is associated with the Brasiliano tectonothermal event that happened around 600 Ma (Dorr II 1969, Morteani *et al.* 2002). The topaz-bearing veins essentially crosscut the altered

phyllites of the Piracicaba Group rocks and the dolomitic marbles of the Gandarela Formation rocks. Because they have a syncline structure, the topaz-bearing veins crosscut the Quadrilátero Ferrífero District rocks according to the following inverse lithological sequence: Sabará and Piracicaba Groups, and the Gandarela Formation of the Itabira Group (Dorr II 1969).

The Antônio Pereira mine is geologically related to the Northeastern flank of the Mariana Anticline and is associated with the Brasiliano tectonothermal event that happened around 600 Ma (Morteani *et al.* 2002, Rojas 2008, Rojas *et al.* 2009). The topaz-bearing veins essentially crosscut the altered rocks of the Cauê and Gandarela Formations of the Itabira Group, and the Cercadinho Formation rocks of the Piracicaba Group. Because it has an anticline structure, the topaz-bearing veins crosscut the Quadrilátero Ferrífero District rocks according to the following normal lithological sequence: Quebra Ossos, Nova Lima, Caraça, Itabira and Piracicaba Groups (Dorr II 1969).

Due to the tropical weathering, the imperial topaz-bearing veins originate from a halo of strongly altered rock in the crosscut rocks, which are intensively altered. This alteration originates from a brownish clay-like material, locally designated by the miners as “borra de café” or coffee grounds (Olsen 1971). Due to this very intense alteration, there is no reliable information about the primary country rock related to the imperial topaz mineralization. Additionally, in both mines, the primary mineral assemblage related to the topaz-bearing veins is composed of kaolinite, quartz,

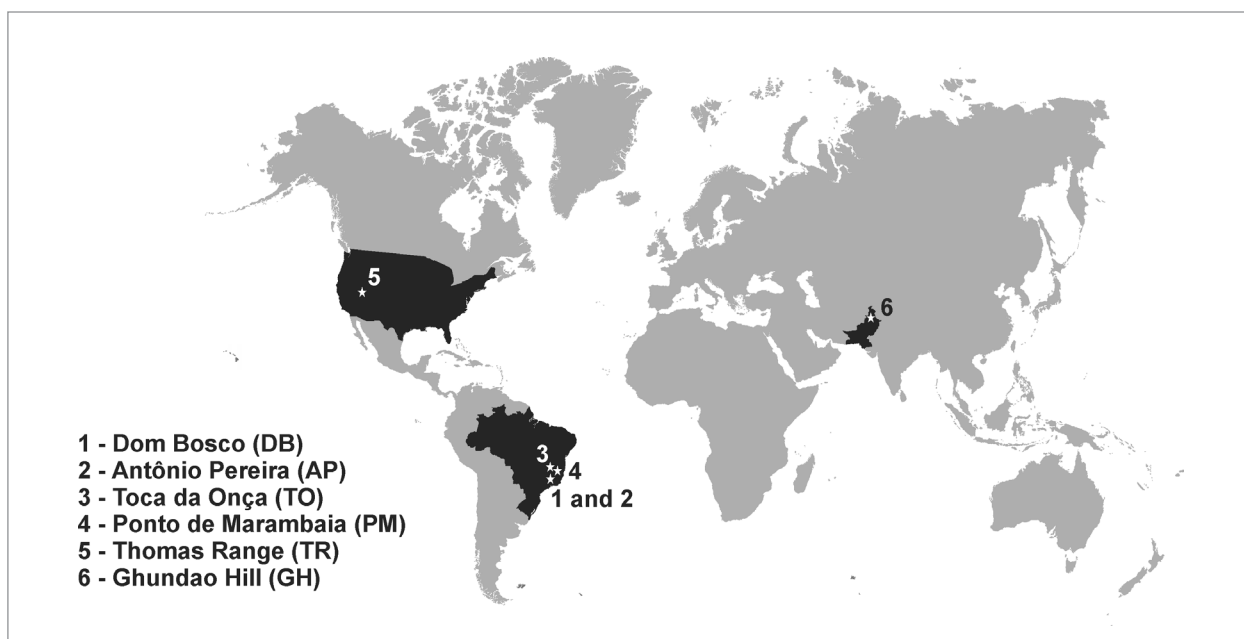


Figure 1. Location of the studied samples: (1) Dom Bosco (DB), (2) Antônio Pereira (AP), (3) Toca da Onça (TO), (4) Ponto de Marambaia (PM); (5) Thomas Range (TR); (6) Ghundao Hill (GH).

muscovite, rutile and, rarely, euclase (Gandini 1994, Morteani *et al.* 2002, Gauzzi *et al.* 2018).

Toca da Onça Pegmatite (Minas Gerais, Brazil)

The Toca da Onça pegmatite is situated in the Northeastern area of the state of Minas Gerais, and is geologically related to the Eastern Brazilian Pegmatite Province. This pegmatite body belongs to the Coronel Murta-Virgem da Lapa Pegmatite Field in the Pegmatite District of Araçuaí. The Toca da Onça pegmatite is intruded in the micaschists of the Salinas Formation (Pedrosa-Soares *et al.* 1990, Morteani *et al.* 2000, Pedrosa-Soares *et al.* 2011). The micaschists of the Salinas Formation are banded and are a primary mineral assemblage composed of predominant biotite, muscovite and plagioclase (albite). Their accessory mineral assemblage is composed of garnet, sillimanite and kyanite (Lima *et al.* 2002). Finally, the Toca da Onça pegmatite body is essentially zoned and complex and its primary mineral assemblage is composed of K-feldspar, quartz and muscovite. The accessory mineral assemblage is composed of tourmaline, cleavelandite, lepidolite, topaz, apatite, columbite-tantalite, amblygonite, spodumene and morganite (Correia-Neves *et al.* 1980, Pinho-Tavares 2004).

Ponto de Marambaia Pegmatite (Minas Gerais, Brazil)

The Ponto de Marambaia pegmatite is situated in the Northeastern area of the state of Minas Gerais, and is geologically related to the Eastern Brazilian Pegmatite Province. This pegmatite belongs to the Padre Paraíso Pegmatite Field in the Pegmatite District of Padre Paraíso. The Ponto de Marambaia pegmatite is specifically intruded in the Caladão granite, which is a porphyritic biotite granitoid, classified as a type-I granite that is metaluminous and K-rich. Its primary mineral assemblage is composed of K-feldspar, quartz, plagioclase, biotite and hornblende. Its accessory mineral assemblage is composed of zircon, apatite, allanite, carbonate, muscovite, chlorite, magnetite, rutile and ilmenite (Ferreira *et al.* 2005). Intruded in the Caladão granite is the Padre Paraíso charnockite, which is a porphyritic biotite granitoid that has a greenish color. Its primary mineral assemblage is composed of K-feldspar, plagioclase, biotite, hornblende and hyperstene. Its accessory mineral assemblage is composed of zircon, apatite, epidote, titanite, carbonate, muscovite, chlorite and magnetite (Ferreira *et al.* 2005). Finally, the Ponto de Marambaia pegmatite body can be zoned or homogeneous, and its primary mineral assemblage is composed of K-feldspar, quartz, biotite, muscovite, beryl (gem-varieties aquamarine, heliodor and goshenite), topaz, schörl and magnetite. The accessory mineral assemblage is composed of sericite, kaolinite, topaz, goethite and Mn-oxides (Ferreira *et al.* 2005).

Thomas Range mine (Utah, United States)

The Thomas Range topaz occurrences are situated in the West-Central part of the state of Utah, and are emplaced in a Cenozoic extensive fault zone, and have ages varying between 0.06–50.00 Ma (Christiansen *et al.* 1986, 2007). Topaz crystals from the Thomas Range occur in topaz rhyolites, as groups of coalesced lava domes and flows with interlayered tephra deposits, whose ages vary between 6.00–7.00 Ma and whose volume is 50 km³ (Christiansen *et al.* 1986, Duffield & Dalrymple 1990). The topaz rhyolites from the Thomas Range mainly show an aphyric texture and are composed 20% by phenocrysts (Lindsey 1982). Their primary mineral assemblage is composed of quartz, plagioclase, K-feldspar, Fe-Ti oxides, biotite, Fe-hornblende, Fe-augite, garnet, and their accessory mineral assemblage is composed of fluorite, pseudobrookite, bixbyite, magnetite, hematite, cassiterite, topaz, ilmenite, zircon, F-titanite, allanite and apatite (Burt *et al.* 1982, Christiansen *et al.* 1986, 2007).

Ghundao Hill mine (Mardan, Pakistan)

The Ghundao Hill topaz mines are situated close to the small town of Katlang (District of Mardan, Pakistan). According to Jan (1979), the topaz-bearing veins were geologically related to the Lower Swat Buner Schistose Group, namely to the unit of Marbles and Calcareous Schists, which was classified by Martin *et al.* (1962). However, after the studies carried out by DiPietro *et al.* (1993), the mentioned group corresponds to the Alpurai Group, which is composed, along with others, the Kashala Formation. The topaz-bearing veins occur in the rocks of the Kashala Formation, which are essentially composed of a sequence of garnet- and/or epidote-bearing calcareous schist that grade into schistose marble and into a coarse grey to brown calcite marble (DiPietro *et al.* 1993). These veins cross-cut the schistosity of the aforementioned marbles and are mainly composed of coarse sparry calcite, whose colors can be white, grey and brownish, quartz and, rarely, laths of muscovite (Jan 1979, Morteani & Voropaev 2007). Finally, according to Morteani and Voropaev (2007), the topaz mineralization resulted from metamorphic fluids, which occurred during an Eocene Himalayan tectonothermal event (43.20 ± 1.70 Ma).

MATERIALS AND METHODS

Samples preparation

According to their respective provenances, the samples were named Dom Bosco (DB), Antônio Pereira (AP), Toca da Onça (TO), Ponto de Marambaia (PM), Thomas Range (TR) and Ghundao Hill (GH). The selected samples showed

a euhedral to subhedral prismatic habit, and had *striae* along the well-preserved lateral faces, which were parallel to the *c*-axis and were free of inclusions (Tab. 1). All of these samples were cut parallel to (001) and perpendicular to the *c*-axis. After this procedure, the samples were mounted in an epoxy resin, and then polished using a 0.25 μm diamond powder until the topaz crystals were exposed. The samples were cleaned with HNO_3 before the analysis in order to avoid contamination.

Analytical methods

Scan electron microscopy-cathodoluminescence (SEM-CL)

Panchromatic SEM-CL images were obtained using a JEOL JSM-6510 equipped with a CL detector, with a wavelength range of 200–800 nm, in the Microanalysis Laboratory of the Department of Geology at the Federal University of Ouro Preto (Brazil). The accelerating voltage was 20 kV.

Laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS)

Trace elements were analyzed by (LA-ICP-MS) using a Thermo-Finnigan Element II, a single collector sector field (SF) ICL1-MS, coupled with a CETAC UV Nd:YAG 213 nm laser with a Helix ablation cell in the Isotopic Geochemistry Laboratory (LOPAG) of the Department of Geology at the Universidade Federal de Ouro Preto (Brazil). The spot size was 40 μm . Helium combined with argon was a carrier gas and ablation was carried out with a pulse rate of 10 Hz and an energy density of 3.60 J/cm². The LA-ICP-MS analyses were punctually performed on the topaz crystals and 20 s of background acquisition was followed by 60 s of ablation. The software GLITTER 4.0 (Macquarie University) was used for data reduction with NIST SRM610 as a primary external standard and Si as the internal standard (using the average SiO_2 content for

each sample). The secondary external standard was NIST SRM612. Limits of detection were about 0.91–3.57 ppm (Li), 222.58–276.19 ppm (Ca), 4.80–5.90 ppm (Ti), 0.04–0.08 ppm (V), 1.60–3.60 ppm (Cr), 0.20–0.40 ppm (Mn), 131.05–300.53 ppm (Fe), 0.06–0.12 ppm (Cu), 0.27–0.50 ppm (Zn), 0.02–0.08 ppm (Nb), 0.08–0.62 ppm (LREE), 0.03–0.12 ppm (HREE), 0.04–0.26 ppm (Ta), 0.07–0.13 ppm (W), and the external reproducibility of the results was 2 to 7%.

RESULTS AND DISCUSSION

Crystallization dynamics

In order to understand the crystallization dynamics of each sample, it was necessary to relate the textural data, obtained by SEM-CL, and the compositional data, obtained by LA-ICP-MS.

To understand the CL textures and their influence on the crystallization dynamics of the minerals according to their geological setting, studies on SEM-CL images, found in the literature, were essential (Boggs *et al.* 2002, Van der Kerkhof *et al.* 2004, Rusk *et al.* 2008, Götte *et al.* 2011, Agangi *et al.* 2014, Frelinger *et al.* 2015, Agangi *et al.* 2016). As observed in Figure 2, samples DB, AP and TR show marked CL-heterogeneity, while samples GH, TO and PM show CL-homogeneity.

The role of Li, Ca, Ti, V, Cr, Mn, Fe, Cu, Zn, Nb, LREE, HREE, Ta and W in the core and rims of topaz crystals was crucial to distinguish the different geological settings (Tab. 2) and to understand their crystallization dynamics. Also, associating the concentrations of these elements with the CL

Table 1. Features of the samples used in this study.

Provenance	Color	Dimensions (width x height; mm)
Dom Bosco (DB)	Orange-pinkish	11.50 × 15.00
Antônio Pereira (AP)	Orangish	7.50 × 4.00
Toca da Onça (TO)	Colorless	9.00 × 12.50
Ponto de Marambaia (PM)	Colorless	14.00 × 12.00
Thomas Range (TR)	Brownish	7.00 × 9.50
Ghundo Hill (GH)	Pinkish	6.00 × 13.00

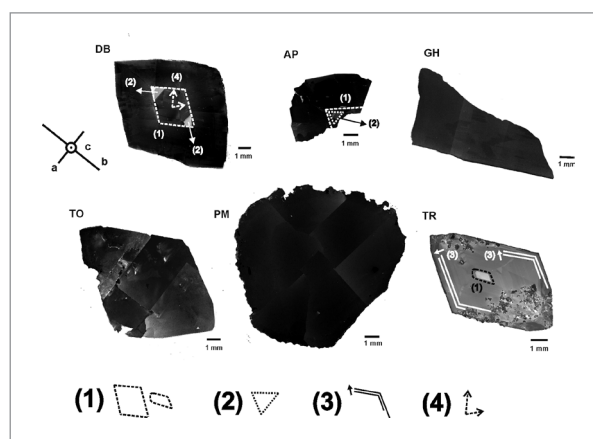


Figure 2. SEM-CL panchromatic images of the different topaz crystals: (1) central rhombic pattern developing along [110]; (2) triangular area with strong luminescence; (3) straight growth zones aligned with [110]; (4) [110] direction.

Table 2. Trace element concentrations obtained in the studied topaz crystals by LA-ICP-MS (in ppm).

Analysis (ppm)	DB 1 Core	DB 2 Core	DB Average Core	DB 3 Rim	DB 4 Rim	DB Average Rim	DB Total Average
Li	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Ca	831.29	5,219.63	3,025.46	480.94	426.96	453.95	1,739.71
Ti	96.51	480.22	288.37	b.d.l.	45.79	45.79	207.51
V	12.84	47.70	30.27	5.57	38.01	21.79	26.03
Cr	252.25	240.33	246.29	230.71	14.47	122.59	184.44
Mn	12.31	15.07	8.69	11.44	2.82	7.13	10.41
Fe	970.93	b.d.l.	970.93	b.d.l.	b.d.l.	b.d.l.	970.93
Cu	2.98	79.46	41.22	3.79	1.53	2.66	21.94
Zn	19.58	59.97	39.78	b.d.l.	6.04	6.04	28.53
Nb	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
LREE			b.d.l.			b.d.l.	b.d.l.
HREE			b.d.l.			b.d.l.	b.d.l.
Ta	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
W	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Analysis (ppm)	AP 1 Core	AP 2 Core	AP Average Core	AP 3 Rim	AP 4 Rim	AP Average Rim	AP Total Average
Li	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Ca	6,801.91	1,384.43	4,093.17	762.40	4,621.05	2,691.73	3,392.45
Ti	b.d.l.	117.47	117.47	12.88	b.d.l.	12.88	65.175
V	31.14	51.34	41.24	43.80	b.d.l.	43.80	42.09
Cr	375.84	1,333.32	854.58	146.5	81.50	114.00	201.28
Mn	b.d.l.	47,00	47,00	3.78	b.d.l.	3.78	25.39
Fe	b.d.l.	955.88	955.88	b.d.l.	2,955.38	2,955.38	1,955.63
Cu	1.93	43.61	22.77	64.22	9.54	36.88	29.825
Zn	9.85	18.66	14.26	29.31	b.d.l.	29.31	19.27
Nb	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
LREE			b.d.l.			b.d.l.	b.d.l.
HREE			b.d.l.			b.d.l.	b.d.l.
Ta	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
W	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Analysis (ppm)	TO 1 Core	TO 2 Core	TO Average Core	TO 3 Rim	TO 4 Rim	TO Average Rim	TO Total Average
Li	9.55	b.d.l.	9.55	b.d.l.	b.d.l.	b.d.l.	9.55
Ca	749.91	604.60	677.26	479.02	b.d.l.	479.02	611.18
Ti	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
V	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Cr	22.11	b.d.l.	22.11	41.21	33.88	37.55	32.40
Mn	1.91	b.d.l.	1.91	8.22	5.94	7.08	5.36
Fe	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Cu	5.85	1.51	3.68	4.58	5.95	5.27	4.47
Zn	5.10	b.d.l.	5.10	b.d.l.	4.72	4.72	4.91
Nb	0.39	4.74	2.57	0.36	0.12	0.24	1.40
LREE			1.00			2.82	1.91
HREE			0.15			b.d.l.	b.d.l.
Ta	b.d.l.	b.d.l.	b.d.l.	2.84	b.d.l.	2.84	2.84
W	b.d.l.	18.93	18.93	b.d.l.	2.18	2.18	10.56

Continue...

Table 2. Continuation.

Analysis (ppm)	PM 1 Core	PM 2 Core	PM Average Core	PM 3 Rim	PM 4 Rim	PM Average Rim	PM Total Average
Li	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Ca	449.08	b.d.l.	449.08	912.4	b.d.l.	912.4	680.74
Ti	48.33	b.d.l.	48.33	20.49	78.26	49.38	49.03
V	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Cr	16.51	b.d.l.	16.51	20.28	48.24	34.26	28.34
Mn	2.99	b.d.l.	2.99	b.d.l.	b.d.l.	b.d.l.	2.99
Fe	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Cu	3.08	b.d.l.	3.08	3.8	b.d.l.	3.8	3.44
Zn	3.33	b.d.l.	3.33	b.d.l.	b.d.l.	b.d.l.	3.33
Nb	b.d.l.	3.04	3.04	b.d.l.	b.d.l.	b.d.l.	3.04
LREE			b.d.l.			b.d.l.	b.d.l.
HREE			1.14			1.42	1.28
Ta	b.d.l.	0.48	0.48	b.d.l.	b.d.l.	b.d.l.	0.48
W	0.49	1.01	0.75	b.d.l.	b.d.l.	b.d.l.	0.75
Analysis (ppm)	TR 1 Core	TR 2 Core	TR Average Core	TR 3 Rim	TR 4 Rim	TR Average Rim	TR Total Average
Li	23.95	b.d.l.	23.95	b.d.l.	13.61	13.61	18.78
Ca	b.d.l.	295.03	295.03	479.00	b.d.l.	479.00	387.02
Ti	46.17	b.d.l.	b.d.l.	11.39	57.00	34.20	38.19
V	b.d.l.	22.76	22.76	b.d.l.	b.d.l.	b.d.l.	22.76
Cr	8.95	5.50	7.23	41.21	b.d.l.	41.21	18.55
Mn	b.d.l.	b.d.l.	b.d.l.	8.22	b.d.l.	8.22	8.22
Fe	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Cu	1.12	2.68	1.90	4.98	b.d.l.	4.98	2.93
Zn	4.34	4.69	4.52	b.d.l.	6.38	6.38	5.14
Nb	77.10	7.47	42.29	0.36	2.79	1.58	21.93
LREE			2.05			b.d.l.	2.05
HREE			b.d.l.			b.d.l.	b.d.l.
Ta	15.51	6.45	10.98	b.d.l.	9.19	9.19	10.38
W	29.99	2.60	16.30	0.18	58.31	29.25	22.77
Analysis (ppm)	GH 1 Core	GH 2 Core	GH Average Core	GH 3 Rim	GH 4 Rim	GH Average Rim	GH Total Average
Li	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Ca	697.80	b.d.l.	697.8	1,192.85	b.d.l.	1,192.85	945.33
Ti	b.d.l.	2.39	b.d.l.	b.d.l.	b.d.l.	b.d.l.	2.39
V	26.81	b.d.l.	26.81	4.62	82.23	43.43	37.89
Cr	61.58	336.58	199.08	80.80	171.46	126.13	162.61
Mn	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Fe	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
Cu	2.56	7.56	5.06	3.65	8.60	6.13	5.59
Zn	b.d.l.	7.41	7.41	8.01	15.37	11.69	10.26
Nb	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
LREE			1.68			3.10	2.39
HREE			b.d.l.			3.16	3.16
Ta	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.
W	b.d.l.	b.d.l.	b.d.l.	2.10	2.30	2.20	2.20

images shows how they are incorporated and whether or not they play a role as hypothetical CL-activators.

In sample DB, SEM-CL images show a central rhombic pattern, which develops along [110], in its core. In the apices of this central rhombic pattern, especially along [010], there is a common tendency to form a triangular area of very strong luminescence (see Fig. 2, DB). This CL-heterogeneity demonstrates that this area was marked by multiple precipitation-dissolution textures, which are representative of physicochemical changes within the hydrothermal system. In the rims, it is possible to observe very dark and incipient luminescence areas, which are truncated by some irregular-shaped areas. These textures can be related to metamorphic activity, shown by lower CL along the rims of the grain, which can imply recrystallization and removal of impurities, thus generating a homogeneous CL response. In addition, all these crystallization features, showed by the CL-panchromatic images, lead one to infer that a primary process of crystallization took place followed by recrystallization process (Gauzzi *et al.* 2018). The sample AP, besides being just part of a rhombic pattern and showing less luminescence, displays the same luminescence features of the sample DB (see Fig. 2, AP).

According to Morteani *et al.* (2002) and Morteani and Voropaev (2007), the crystallization dynamics of samples

DB and AP could be related to the metamorphic tectono-thermal event that occurred during the Brasiliano orogeny in the region of Ouro Preto in about 600 Ma. In comparison with the study carried out by Gauzzi *et al.* (2018) on imperial topaz crystals related to the Dom Bosco mine, the same CL behavior was observed as in the samples of our study (Figs. 3A and 3B). According to the chemical composition data, in both samples, the high Ca, Ti, Cr and Mn concentrations are observable in the cores (see Tabs. 2 and 3, and Figs. 3C, 3D and 3E). However, V, Fe, Cu and Zn concentrations are higher in the core than in the rims, namely in sample DB (Fig. 3D), which is corroborated by the results obtained by Gauzzi *et al.* (2018) and in Table 3. Meanwhile, the opposite can be observed in sample AP (Fig. 3E).

Considering that crystallization dynamics can be related to primary crystallization and recrystallization processes in imperial topaz, which is corroborated by the SEM-CL images and by Gauzzi *et al.* (2018), the incorporation of Ca, Ti, Cr and Mn may have been very intense during the first stages of crystallization in both samples. On the other hand, the incorporation of V, Fe, Cu and Zn allows for both samples to be distinguished from one another. In the sample DB, these elements were intensively aggregated to its crystal structure during the first stages of crystallization, while the opposite occurred in sample AP. Also, the timing of incorporation

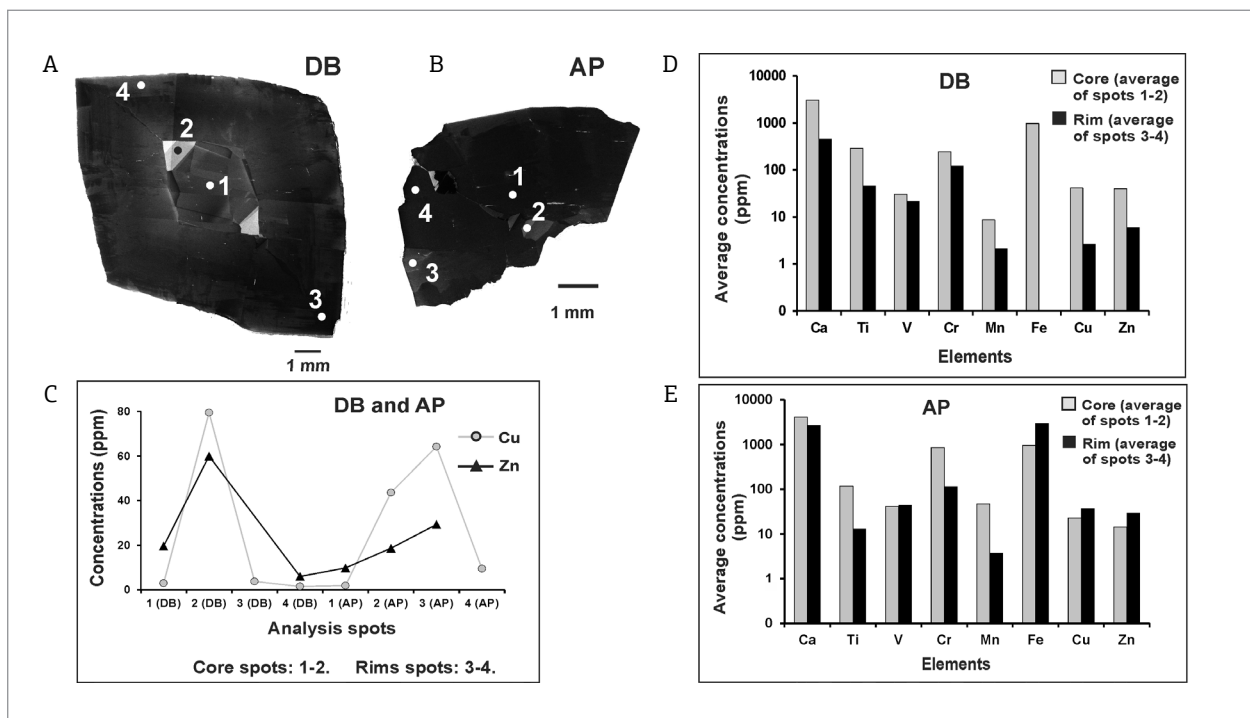


Figure 3. A combination of the SEM-CL image and the chemical composition in samples DB and AP: (A) and (B) SEM-CL images with the LA-ICP-MS spots in samples DB and AP, respectively; (C) the relation between the Cu and Zn concentrations and the CL-intensities observed in (A) and (B); (D) and (E) concentrations of Ca, Ti, V, Cr, Mn, Fe, Cu and Zn (in ppm) in the core and in the rims of both samples, respectively.

Table 3. A comparison between the trace element composition of topaz crystals from our study (i.e. DB) and from different geological environments found in the literature (average values in ppm).

	Geological setting	Ca (average; ppm)		Ti (average)		V (average)		Cr (average)		Mn (average)	
		Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim
Agangi <i>et al.</i> (2014, 2016)	Ongonite	588.50	1,293.05	18.32	30.22	0.27	1.47	b.d.l.	b.d.l.	7.68	24.10
Gauzzi <i>et al.</i> (2018)	Hydrothermal veins	n.a.	n.a.	215.74	161.92	28.71	23.70	319.42	300.87	5.09	3.65
Our study											
DB	Hydrothermal veins	3,025.46	453.95	288.37	45.79	30.27	21.79	246.29	12.59	8.69	2.13
AP	Hydrothermal veins	4,093.17	2,691.73	117.47	12.88	41.24	40.80	854.58	114.00	47.00	3.78
TO	Pegmatite	677.26	479.02	b.d.l.	b.d.l.	b.d.l.	b.d.l.	22.11	37.55	1.91	7.08
PM	Pegmatite	449.08	912.40	48.33	49.38	b.d.l.	b.d.l.	16.51	34.26	2.99	b.d.l.
TR	Rhyolite	295.03	479.00	43.75	34.20	22.76	b.d.l.	7.23	41.21	b.d.l.	8.22
GH	Hydrothermal veins (metamorphism)	697.80	1,192.85	b.d.l.	b.d.l.	26.81	43.43	199.08	126.13	b.d.l.	b.d.l.
	Geological setting	Fe (average; ppm)		Cu (average)		Zn (average)		Nb (average)		LREE (average)	
		Core	Rim	Core	Rim	Core	Rim	Core	Rim	Core	Rim
Agangi <i>et al.</i> (2014, 2016)	Ongonite	60.64	132.97	2.00	10.84	1.90	13.98	1.36	20.14	0.50	2.63
Gauzzi <i>et al.</i> (2018)	Hydrothermal veins	799.07	691.73	13.73	10.35	19.75	11.33	n.a.	n.a.	n.a.	n.a.
Our study											
DB	Hydrothermal veins	970.93	b.d.l.	41.22	2.66	24.78	6.04	b.d.l.	b.d.l.	b.d.l.	b.d.l.
AP	Hydrothermal veins	955.88	2,955.38	22.77	36.88	14.26	29.31	b.d.l.	b.d.l.	b.d.l.	b.d.l.
TO	Pegmatite	b.d.l.	b.d.l.	3.68	5.27	5.10	4.72	2.57	0.24	1.00	2.82
PM	Pegmatite	b.d.l.	b.d.l.	3.08	3.80	3.33	b.d.l.	3.04	b.d.l.	b.d.l.	b.d.l.
TR	Rhyolite	b.d.l.	b.d.l.	1.90	4.98	4.52	6.38	42.29	1.58	2.05	b.d.l.
GH	Hydrothermal veins (metamorphism)	b.d.l.	b.d.l.	5.06	6.13	7.41	11.69	b.d.l.	b.d.l.	1.68	2.39
	Geological setting	HREE (average; ppm)		Ta (average)		W (average)					
		Core	Rim	Core	Rim	Core	Rim				
Agangi <i>et al.</i> (2014, 2016)	Ongonite	0.03	0.74	0.64	3.73	1.69	18.42				
Gauzzi <i>et al.</i> (2018)	Hydrothermal veins	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.				
Our study											
DB	Hydrothermal veins	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.				
AP	Hydrothermal veins	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.	b.d.l.				
TO	Pegmatite	0.15	b.d.l.	b.d.l.	2.84	18.93	2.18				
PM	Pegmatite	1.14	1.42	0.48	b.d.l.	0.75	b.d.l.				
TR	Rhyolite	b.d.l.	b.d.l.	5.98	9.19	16.30	29.25				
GH	Hydrothermal veins (metamorphism)	b.d.l.	3.16	b.d.l.	b.d.l.	b.d.l.	2.20				

n.a. (not analyzed); b.d.l. (below detection limit)

of the trace elements can be observed in the CL intensity. In agreement with Gauzzi *et al.* (2018), Cu and Zn could be considered important CL-activators in imperial topaz from the Ouro Preto region and, thus, Figures 3A, 3B and 3C show that high concentrations of these elements are related to high-CL areas. In this case, because the incorporation of these elements has been much more significant in the first stages of crystallization of sample DB, due to their high concentrations in the core (Fig. 3D), the CL intensity is much more marked and evident (Fig. 3A). Notwithstanding, in sample AP, Cu and Zn must have been incorporated in its last stages of crystallization, due to their high concentrations in the rims (Fig. 3E), and thus the CL intensity is more discrete (Fig. 3B) than the one observed in sample DB.

Both samples TO and PM show low luminescence areas and are predominantly dark and completely homogeneous (Fig. 2, TO and PM). In both samples, no difference was noticed between the core and the rim in terms of luminescence. This can imply a slow and late crystallization, typical of late-stage minerals that crystallize in pegmatites. Moreover, this homogeneity allows for the reduction of structural defects, the removal of impurities of the crystal structure, and a homogeneous CL response to be observed.

In agreement with the chemical composition data, in sample TO, higher Ca, Zn, Nb, HREE and W concentrations are more observable in the core than in the rims (Tabs. 2 and 3, and Figs. 4A and 4B). However, the opposite can be observed with elements such as Cr, Mn, Cu, LREE and Ta. On the other hand, in sample PM, higher Mn, Zn, Nb, Ta and W concentrations are more evident in the core than in the rims, while the opposite can be noticed with elements such as Ca, Ti, Cr, Cu and HREE (Tabs. 2 and 3, and Figs. 4C and 4D).

The crystallization dynamics can be very similar in these samples due to the slow crystallization processes, and the dark and homogeneous images obtained by SEM-CL. Nonetheless, the chemical composition is a very relevant distinguishing aspect. The incorporation of Li, Zn, Nb, HREE and W may have been very intense during the first stages of crystallization in sample TO, while elements such as Cr, Mn, Cu, LREE and Ta were aggregated to its crystal structure during the final stages (Fig. 4B). On the other hand, in sample PM, elements such as Mn, Zn, Nb, Ta and W may have been strongly incorporated during its first stages of crystallization, while Ca, Ti, Cr, Cu and HREE were more likely to enter its crystal structure during the final

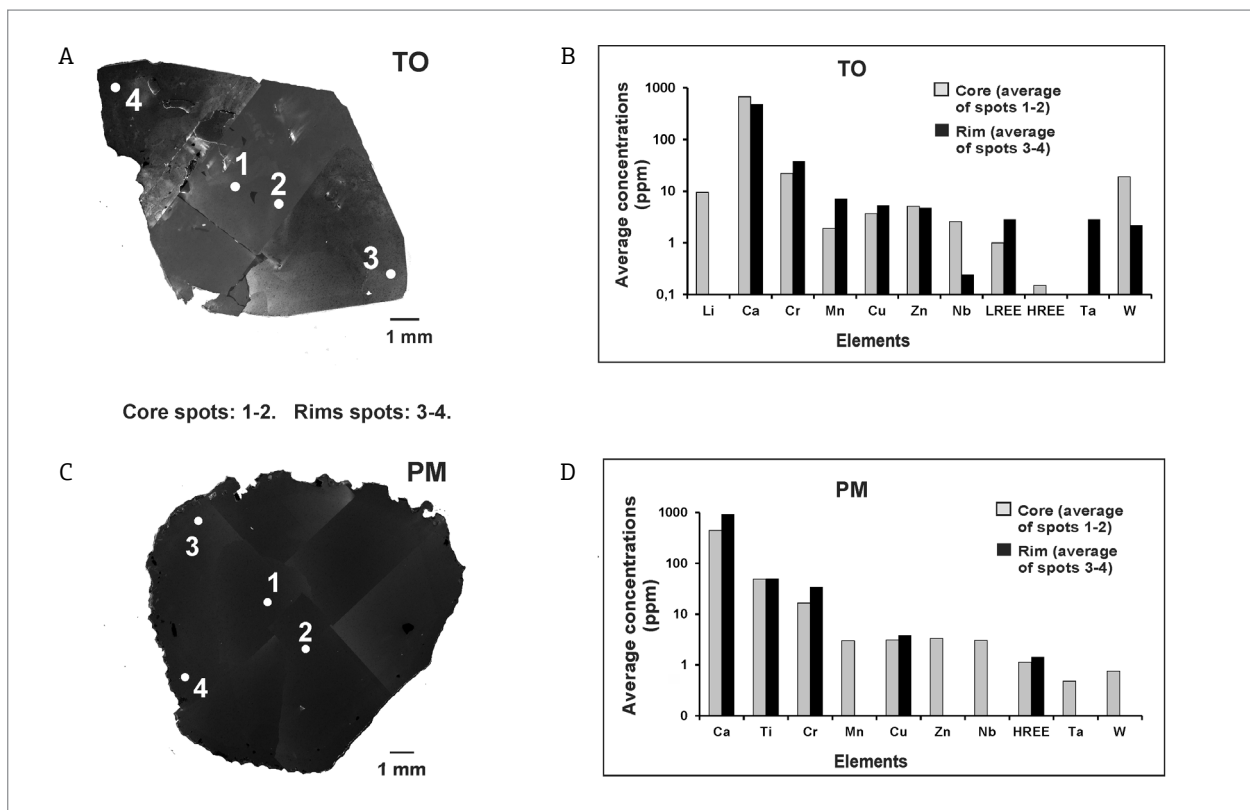


Figure 4. A combination of the SEM-CL image and the chemical composition in samples TO and PM: (A) and (C) SEM-CL images with the LA-ICP-MS spots; concentrations of (B) Li, Ca, Cr, Mn, Cu, Zn, Nb, LREE, HREE, Ta and W, and of (D) Ca, Ti, Cr, Mn, Cu, Zn, Nb, HREE, Ta and W (both in ppm) in the core and in the rims.

stages (Fig. 4D). Moreover, as both samples show dark and homogeneous CL intensities (Figs. 4A and 4C), no relation between the concentrations of these trace elements and CL-activation has been found.

Sample TR displays high CL intensities. In its core, it is possible to observe a growth sector boundary, specifically a very high luminescence central rhombic pattern developing along [110]. In the rims, it is possible to observe some straight growth zones with oscillating, but discrete, CL intensities, which are aligned with [110] (Fig. 2, TR). In magmatic topaz, it is possible to observe well-preserved growth zones of oscillating CL intensities that can be truncated by high luminescence irregular areas. This can be explained by the periods of crystal growth, resorption and the subsequent growth again of late-stage topaz in rhyolitic magmas of the Thomas Range (Christiansen *et al.* 1986, 2007). Sample TR can be compared to the ongonite topaz samples, studied by Agangi *et al.* (2014, 2016), due to the fact that it is a magmatic topaz (ongonite from Ary-Bulak Massif, Russia). The SEM-CL image of sample TR shows very high luminescence in the core and well-defined growth zones according to [110]. Also, the weak luminescence areas show a marked grey that is not dark. As observed in the studies carried out by Agangi *et al.* (2014, 2016), truncated irregular areas are observable. However, while the topaz samples of the latter

studies show CL-variations that define euhedral to lobed growth textures, the sample TR only shows euhedral and growth textures.

In sample TR, the chemical composition data allow for higher Li, V, Nb, LREE and Ta concentrations to be observed in the core than in the rims, while the opposite can be noticed with elements such as Ca, Ti, Cr, Mn, Cu, Zn and W (Tabs. 2 and 3, and Figs. 5A, 5B and 5C).

In agreement with Figure 5C and Agangi *et al.* (2014, 2016), the crystallization dynamics may have been marked by a strong incorporation of Li, V, Nb, LREE and Ta during its first stages of crystallization. Meanwhile, Ca, Ti, Cr, Mn, Cu, Zn and W were aggregated to its crystal structure during the final stages (Fig. 5C). In agreement with Agangi *et al.* (2014, 2016), Li, Nb, Ta and W could be considered important CL-activators in magmatic topaz and, thus, Figures 5A and 5B show that high concentrations of these elements are related to high intensity CL areas. Moreover, because of the high concentrations of these CL-activator elements, above tens of ppm (Fig. 5C), the SEM-CL image shows high CL intensity and an absence of dark areas.

Sample GH shows weak and dull luminescence areas, and homogeneous CL response (Fig. 2, GH). This is related to metamorphism, especially to recrystallization, which increases the crystal structure order, reducing the structural

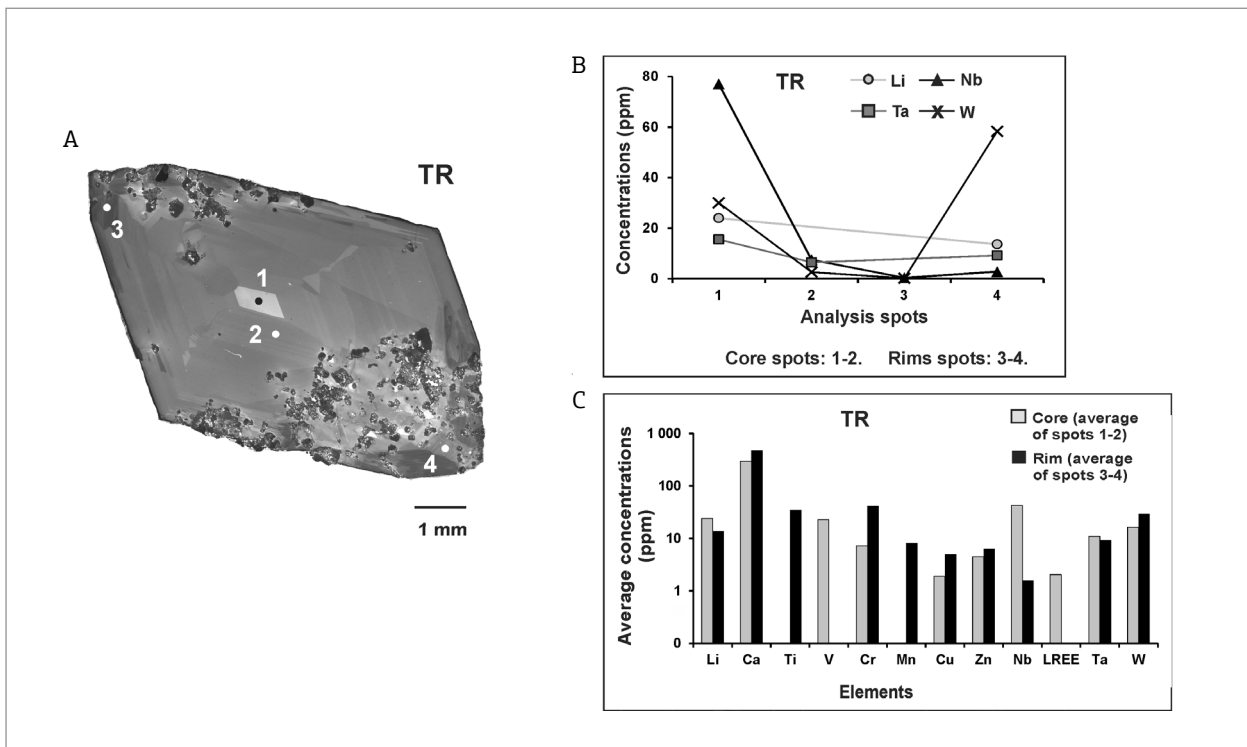


Figure 5. A combination of the SEM-CL image and the chemical composition in sample TR: (A) a SEM-CL image with the LA-ICP-MS spots; (B) the relation between the Li, Nb, Ta and W concentrations and the CL-intensities observed in (A); (C) concentrations of Li, Ca, Ti, V, Cr, Mn, Cu, Zn, Nb, LREE, Ta and W (in ppm) in the core and in the rims.

defects in the crystal and removing the impurities from the crystal structure. Also, according to Morteani and Voropaev (2007), topaz crystallization is related to the metamorphic fluids that originated from the Eocene Himalayan tectono-thermal event (43.20 ± 1.70 Ma).

The chemical composition data, in sample GH, show higher Ti and Cr concentrations in the core than in the rims, while the opposite can be observed with elements such as Ca, V, Cu, Zn, LREE, HREE and W (Tabs. 2 and 4, and Figs. 6A and 6B).

According to Figure 6B, the crystallization dynamics may have been marked by an intense incorporation of Ti and Cr during the first stages of crystallization, while elements such as Ca, V, Cu, Zn, LREE, HREE and W were aggregated to its crystal structure during the final stages. Furthermore, this sample shows weak and homogeneous CL intensity (Fig. 6A), and the relation between element concentrations and CL-activation is the same as in samples TO and PM.

Table 4. Concentrations of Ti, V, Cr, Mn, Cu and Zn as trace elements in the BIF Lake Superior-type of the Quadrilátero Ferrífero District (Selmi *et al.* 2009, Adekoya *et al.* 2012).

Elements	BIF Lake Superior-type (ppm)
Ti	160–1,900
V	30–124
Cr	102–122
Mn	3,400–4,600
Cu	10–37
Zn	2–42

The influence of primary and accessory mineral assemblages in the fingerprint of the studied topaz crystals

The Li, V, Cr, Mn, Cu, Zn, Nb, LREE, HREE, Ta and W concentrations can occur as part of the chemical formula, or as relevant trace elements of the minerals that constitute the primary and/or accessory mineral assemblages of the rocks in which the studied topaz crystals occurred. Ti concentrations can be related to primary/accessory mineral assemblages (*e.g.* rutile). The Ca and Fe concentrations can be associated with the enrichment of these elements in some of the lithologies (*e.g.* Fe in BIF). Moreover, the concentrations and the combinations of these elements are essential to characterize the fingerprint of each topaz crystal of this study.

The chemical composition data show that in sample DB, high Ca concentrations should have resulted from the topaz mineralizing fluid crosscutting the following sequences: dolomitic marbles, dolomitic phyllites and dolomitic itabirites of the Gandarela Formation in addition to carbonate-bearing phyllites and dolomitic marbles of the Piracicaba Group. And in sample AP, the following sequences: carbonate-bearing phyllites and dolomitic marbles of the Piracicaba Group in addition to dolomitic marbles, dolomitic phyllites, and dolomitic itabirites of the Gandarela Formation. Furthermore, Ca abundance in the mineralizing fluids can be explained, not only by the carbonate host rocks, but also, according to Gandini (1994), by the presence of dolomite as a mineral assemblage and the presence of calcite crystals as mineral microinclusions. The high Ti, V, Cr, Mn, Fe, Cu and Zn concentrations in sample AP, can be related to the mineralizing fluid of imperial topaz crosscutting the iron ore bodies of BIF Lake Superior-type rocks of the Cauê

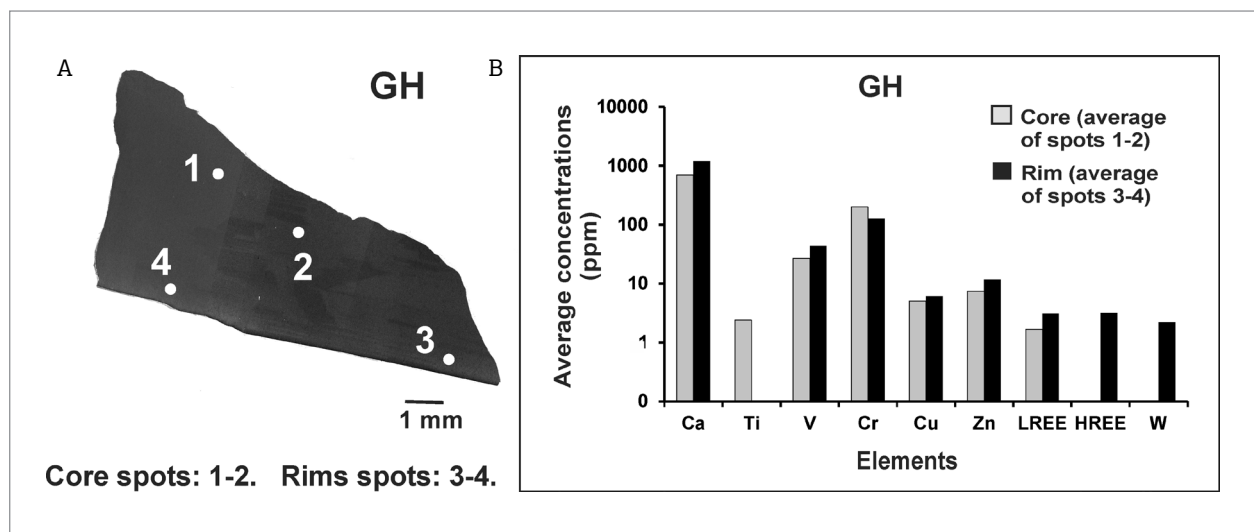


Figure 6. A combination of the SEM-CL image and the chemical composition in sample GH: (A) a SEM-CL image with the LA-ICP-MS spots; (B) concentrations of Ca, Ti, V, Cr, Cu, Zn, LREE, HREE and W (in ppm) in the core and in the rims.

Formation (sample AP). On the other hand, the concentrations of these elements in sample DB can be related to the dolomitic itabirites, which are Fe-rich (FeO concentrations of up to 42.50 wt.%; Spier *et al.* 2003), of the Gandarela Formation. In both cases, the percolation and capture of Fe by the mineralizing fluid explains how the topaz originated. In the case of Ti, this enrichment can be related to these iron ores that originate from sediments, which, according to Correns (1978), are Ti-rich. Cr tends to accumulate in the sedimentary Fe- and Al-rich rocks (Shiraki 1978), and has a high affinity with Fe-oxides (Yu *et al.* 2001). Mn, according to Baioumy *et al.* (2013), has a strong chemical affinity with iron ore deposits, especially in this case with the occurrence of Mn-oxides in these iron deposits (Gandini 1994). Also, the fluid should be Fe-, Cr-, Mn- and Ti-rich, not only because of the crosscutting of the aforementioned BIF rocks and the capturing of their respective Fe and Ti, but also because of the Mn-oxide, hematite, and rutile crystals existent in the main mineral assemblage. The Cr enrichment can be related to the Fe-oxides, specifically hematite and magnetite, and to the already mentioned affinity between Cr and Fe, due to the fact that there is no Cr-bearing mineral in the main mineral assemblage. High V, Cu and Zn concentrations can be related to their presence as trace elements in banded iron formations, which could have been captured during imperial topaz crystallization. In the case of BIF Lake Superior-type rocks, the Ti, V, Cr, Mn, Cu and Zn concentrations range widely (Tab. 4). The concentrations of Ca, V, Cr, Mn, Fe, Cu and Zn are higher in sample AP than in sample DB (Tab. 2 and Fig. 7). Due to the

influence of the mineral assemblages (*e.g.* rutile) and the rocks of the surrounding geological setting (*e.g.* carbonate rocks), the fingerprint of both samples is characterized by Ca, Ti, V, Cr, Mn, Cu and Zn (Fig. 7).

In sample TO, Cr, Cu, Zn and HREE mainly occur as relevant trace elements in the minerals of the primary and accessory mineral assemblages, such as biotite, garnet, pyroxene and amphibole (Herrmann 1978, Shiraki 1978, Wedepohl 1978a, 1978d) in the Salinas Formation rocks. The high Ca concentrations can be related to apatite, plagioclase and titanite, which are part of the primary and accessory mineral assemblages of the Salinas Formation rocks. The high Li concentrations are related to the lepidolite, amblygonite and spodumene, which are part of the accessory mineral assemblage of this pegmatite. The relatively low Mn, Nb and Ta concentrations can be related to the main incorporation of these elements into crystallize columbite-tantalite (Wedepohl 1978b, 1978c, Heinrich 1978). The LREE concentrations can be associated with the crystallization of allanite. Also, the Mn concentrations are relevant enough to act as a chromophore element of morganite. The W concentrations, according to Krauskopf (1978), are associated with the Nb-Ta-bearing minerals such as columbite-tantalite, whose W concentrations can reach up to 1.00 wt.%, due also to the substitution of W^{6+} for Nb^{5+} and Ta^{5+} . Also, the W concentrations can be associated with the muscovite. According to the influence of the mineral assemblages (*e.g.* columbite-tantalite) and to the rocks of the surrounding geological setting (*e.g.* pegmatites), the fingerprint of this sample is characterized by Li, Nb, LREE, Ta and W (see Fig. 7).

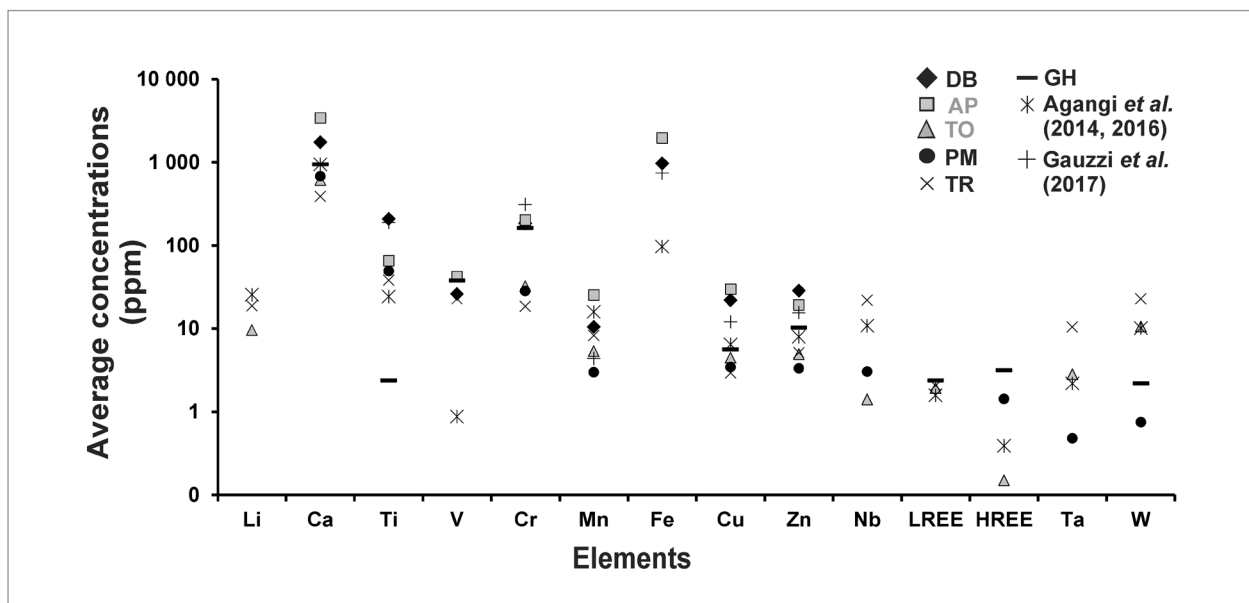


Figure 7. A comparison between the concentrations of Li, Ca, Ti, V, Cr, Mn, Fe, Cu, Zn, Nb, LREE, HREE, Ta and W in the studied samples and in the samples of Agangi *et al.* (2014, 2016) and Gauzzi *et al.* (2018).

In sample PM, the Ti and Cr concentrations could be related to the Fe-oxides, especially the goethite. The Mn concentrations could be related to the Mn-oxides. Both Fe- and Mn-oxides occur as part of the accessory mineral assemblage of this pegmatite. The Ti concentrations could also be related to the rutile and ilmenite, which occur as part of the accessory mineral assemblage of the Caladão Granite, and to the titanite, which occurs as part of the accessory mineral assemblage of the Padre Paraíso Chernockite. Ca, Cu, Zn, Nb, HREE, Ta and W mainly occur as relevant trace elements in the minerals of the following primary and accessory mineral assemblages of the Caladão Granite. Ca concentrations are related to apatite, allanite and epidote. Cu and HREE are related to biotite, hornblende, hyperstene and plagioclase (Herrmann 1978, Wedepohl 1978a). Zn is related to magnetite and muscovite (Wedepohl 1978d). Nb is related to hyperstene and titanite (Heinrich 1978). W is related to muscovite (Krauskopf 1978). The mineral assemblages (*e.g.* titanite) and the rocks of the surrounding geological setting (*e.g.* pegmatites) were very relevant for the fingerprint of this sample, characterized by Ti, Nb and HREE (Fig. 7).

In sample TR, the low Ca and high Li concentrations could be related to the low CaO (below 0.80 wt.%; Christiansen *et al.* 1984, 1986) and to the Li concentrations (between 37.00–60.00 ppm) in the rhyolites from the Thomas Range. Part of this Ca should have been incorporated in the crystallization of Ca-bearing minerals such as apatite, fluorite and F-titanite, which are among the accessory mineral assemblage of these rhyolitic rocks. Ti, Cr and Mn concentrations can be related to TiO₂ (average 0.10 wt.%), MnO (average 0.06 wt.%) and Cr (average 2.00 ppm) concentrations in the rhyolites from the Thomas Range (Christiansen *et al.* 1986). Part of this Ti and Mn should have been used in the crystallization of the primary (Fe-oxides) and accessory (ilmenite and F-titanite) mineral assemblage that occurs with topaz in the rhyolitic rocks of the Thomas Range. The relatively high Ti concentrations could be related to Ti-bearing minerals such as ilmenite, Fe-Ti oxides, pseudobrookite and F-titanite, which are Ti-bearing minerals. Cr and Mn have a high affinity with Fe-oxides (Yu *et al.* 2001), such as hematite and magnetite, and with the minerals of the primary mineral assemblage such as Fe-augite and Fe-hornblende. V, Cu, Zn, and W mainly occur as relevant trace elements in the minerals of the primary and accessory mineral assemblages such as apatite, biotite, F-titanite, Fe-augite, Fe-hornblende, Fe-Ti oxides and ilmenite (Landergrén 1978, Wedepohl 1978a, 1978d, Krauskopf 1978). Concentrations of Nb (between 53.00–64.00 ppm), LREE (between 7.50–73.00 ppm) and Ta (between 2.30–6.40 ppm) could be related to Nb, LREE and Ta concentrations in the rhyolites from the Thomas

Range. The LREE concentrations could be associated with the main incorporation of these elements to crystallize allanite (Christiansen *et al.* 1984, 1986). The mineral assemblages (*e.g.* ilmenite) and the surrounding rhyolitic rocks influenced the fingerprint of this sample, which is characterized by Li, Ti, V, Mn, Nb, LREE, Ta and W (Fig. 7).

In sample GH, the high Ca concentrations are also directly related to the topaz mineralizing fluid that cross-cuts the following carbonate rocks: calcareous schists, schistose marbles, calcite marbles, and dolomitic marbles of the Kashala Formation. The Zn and W concentrations, according to Wedepohl (1978d) and Krauskopf (1978), can be associated with the muscovite that occurs in the topaz-bearing veins, as part of their mineral assemblage with quartz and topaz crystals. Elements such as V, Cr, Cu, LREE and HREE mainly occur as relevant trace elements in the minerals of the primary mineral assemblages of the rocks of the Kashala Formation such as garnet, ilmenite, muscovite and plagioclase (Herrmann 1978, Landergrén 1978, Shiraki 1978, Wedepohl 1978a). In accordance with the influence of the mineral assemblages (*e.g.* muscovite) and of the surrounding carbonate rocks, the fingerprint of this sample is characterized by Ca, Cr, V, Zn, LREE, HREE and W (Fig. 7).

CONCLUSIONS

Samples DB and AP allow for primary crystallization and recrystallization processes to be observed. They are probably associated with the metamorphic tectonothermal event that occurred during the Brasiliano orogeny, about 600 Ma. This is suggested by the SEM-CL images, in which the cores are CL heterogeneous, and the rims are dark and homogeneous. In both cases, their crystallization dynamics may have been marked by an intense incorporation of Ca, Ti, Cr and Mn during the first stages of crystallization. The aggregation of V, Fe, Cu and Zn in the crystal structure allows both samples to be distinguished. In sample DB, this is evident in the first stages of crystallization, while the opposite is observable in sample AP. In both cases, Cu and Zn could play an important role as CL-activators. However the difference of timing and in intensity of how these elements are incorporated could distinctly highlight the brightness of CL intensities. The high CL central area is more evident in sample DB than AP.

In samples TO and PM, their SEM-CL images show homogeneous and dark CL intensities, which suggest slow and late crystallization processes, typical of late-stage minerals that crystallize in pegmatites. However their crystallization dynamics, in terms of the incorporation of trace elements, allows them to be distinguished from one another.

In sample TO, the first stages of crystallization may have been marked by an intense incorporation of Li, Zn, Nb, HREE and W. Nonetheless, Cr, Mn, Cu, LREE and Ta are strongly aggregated in the final stages. On the other hand, sample PM shows the first stages of crystallization marked by the incorporation of significant Mn, Zn, Nb, Ta and W. Nevertheless, in the final stages, Ca, Ti, Cr, Cu and HREE enter the crystal structure intensively. Even though both samples show dark and homogeneous CL intensities, no relation between the concentrations of these trace elements and CL-activation has been found.

The SEM-CL images of sample TR display well preserved growth zones of oscillating CL intensities, suggesting periods of crystal growth, resorption, and subsequent growth of late-stage topaz in rhyolitic magmas. The crystallization dynamics of this sample may have been marked by a strong incorporation of Li, V, Nb, LREE and Ta during its first stages of crystallization. However, elements such as Ca, Ti, Cr, Mn, Cu, Zn and W have been aggregated to its crystal structure during the final stages. Li, Nb, Ta and W can be considered important CL-activators in magmatic topaz and thus their high concentrations allow for high CL intensity and a lack of dark areas to be observed in the SEM-CL image.

Sample GH allows for a recrystallization process to be observed. It is probably related to the metamorphic activity that occurred during the Eocene Himalayan tectonothermal event (43.20 ± 1.70 Ma). This is suggested by the SEM-CL images, in which the topaz shows weak, dull and homogeneous luminescence areas. The crystallization dynamics may have been marked by a strong incorporation of Ti and Cr during the first stages of crystallization, while elements such as Ca, V, Cu, Zn, LREE, HREE and W may have entered

its crystal structure during the final stages. Similar to samples TO and PM, this sample shows weak and homogeneous CL intensity, which prevents making a relation between the trace element concentrations and CL-activation.

Finally, the fingerprint of each studied topaz crystal can be characterized by the combination of the following elements:

- Ca, Ti, V, Cr, Mn, Cu and Zn, for samples DB and AP;
- Li, Nb, LREE, Ta and W, for sample TO;
- Ti, Nb and HREE, for sample PM;
- Li, Ti, V, Mn, Nb, LREE, Ta and W, for sample TR;
- Ca, Cr, V, Zn, LREE, HREE and W, for sample GH.

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