

Petrography and geochronology of the Furquim Quartzite, an eastern extension of the Itacolomi Group (Quadrilátero Ferrífero, Minas Gerais)

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Abstract

This paper presents the results of a petrographic and geochronological investigation of the Furquim Quartzite (FQ) to establish its stratigraphic correlation to quartzitic units of the Quadrilátero Ferrífero (QF) province. The Quartzite comprises a ca. 20km long and 1-6km wide ridge overlying discordantly the Archean to Paleoproterozoic gneissic basement and rocks of the Archean Rio das Velhas Supergroup between the city Mariana and the town Furquim, southeast of the QF. Despite the discordant contacts, previous field-based stratigraphic studies considered the Furquim Quartzite as part of the Archean Maquiné Group – top unit of the Rio das Velhas Supergroup. U-Pb zircon geochronology via LA-ICP-MS identified several detrital populations ranging from Paleoproterozoic to Archean age. The youngest population of 2087 ± 19 Ma defines the maximum age for the sedimentation of the precursor sandstone. This age can be correlated to be the age of the youngest zircon population of the Itacolomi Group quartzites in the QF. Thus, in contrast to previous studies, the results indicate that the FQ is an eastern extension of the Itacolomi Group, the youngest unit of the Paleoproterozoic Minas Supergroup.

Keywords: Furquim Quartzite, Quadrilátero Ferrífero, U/Pb Geochronology, LA-ICP-MS, Paleoproterozoic.

1. Introduction

The Furquim Quartzite is exposed along a ca. 20 km long and 1 to 6 km wide ridge some 10 to 20 km southeast of the Quadrilátero Ferrífero (Fig. 1a). In the investigated area (Fig. 1b) the Furquim Quartzite trends NE-SW and inflects to a NNE-trend in the region of the town of Furquim. According to the geological map of Baltazar *et al.* (1993), the quartzite ridge separates the western

Archean Santa Bárbara Complex from the eastern Proterozoic Mantiqueira Complex (Fig.1a). Baltazar *et al.* (1993) correlated this quartzite unit to the Maquiné Group of the Archean Rio das Velhas Supergroup. However, the postulated correlation with the Rio das Velhas quartzite is not corroborated by the mode of occurrence, since in a great part of the area, the quartzite ridge overlies

discordantly the Rio das Velhas schists and the Santa Bárbara gneisses, thus suggesting an allochthonous origin. This paper presents the results of the petrographic and geochronological investigation of the Furquim Quartzite in order to enable the discussion of its stratigraphic correlation to quartzite units of the well-known Maquiné, Moeda and Itacolomi quartzites in the Quadrilátero Ferrífero.

2. Geological context

The main quartzitic units in the region of the Quadrilátero Ferrífero belong to the Archean Rio das Velhas Supergroup and to the Paleoproterozoic Minas Supergroup.

The Rio das Velhas Supergroup is subdivided into Nova Lima and Maquiné groups (Dorr, 1969). The Nova Lima

Group is composed mostly of metaultramafic, metamafic and metasedimentary pelitic to ruditic rocks. Felsic volcanic rocks mark the final deposition of the Nova Lima Group at ca. 2.75 Ga (Machado *et al.*, 1992, 1996; Noce *et al.*, 2005). The overlying Maquiné Group is a clastic unit comprised of mainly quartz-

ites. U-Pb age determinations of detrital zircons indicate 3.2 to 2.9 Ga for the main sources of the Maquiné sediments (Machado *et al.*, 1996).

The Minas Supergroup overlies the Rio das Velhas Supergroup and surrounding TTG-gneiss terrains. From bottom to top, it is subdivided into the Tamanduá,

Caraça, Itabira, Piracicaba, Sabará, and Itacolomi Groups (Dorr, 1969). Zircon U–Pb detrital age data suggest that the maximum age of deposition of the sediments of the Caraça quartzite is *ca.* 2.6 Ga (Machado *et al.*, 1996; Hartmann *et al.*, 2006). The Sabará and Itacolomi groups are the youngest units of the Minas Supergroup. The Sabará Group comprises metasedimentary rocks such as metadiamicrites and metatuffs and metaturbidites. The Itacolomi Group comprises quartzites derived from sediments with a maximum deposition age of around 2.1 Ga (Machado *et al.*, 1996, Hartmann *et al.*, 2006). Similar ages were also obtained

for the deposition of the Sabará sediments (Machado *et al.*, 1996). Table 1 presents the ages of the youngest zircons found in the main quartzitic units of the Quadrilátero Ferrífero. These ages correspond to the maximum age for the deposition of the sediments.

Two other geological units that would be potential sources for the detrital zircons of the Furquim Quartzite are the TTG–gneiss complexes, including the Santa Bárbara and the Mantiqueira complexes (Fig. 1b). The Santa Bárbara Complex corresponds to an Archean TTG gneiss terrain considered to be the basement of the Rio das Velhas Supergroup

in the eastern portion of the Quadrilátero Ferrífero. Geochronological U–Pb SHRIMP and LA-ICP-MS dating by Lana *et al.* (2013) indicate crystallization ages of 3.2 Ga. The Mantiqueira Complex is composed of TTG orthogneisses thrust over the southern margin of the São Francisco Craton during the 2.1 Ga Transamazonian event (Silva *et al.*, 2002, Noce *et al.*, 2007). U–Pb zircon age determinations by SHRIMP (Silva *et al.*, 2002; Noce *et al.*, 2007) resulted in paleoproterozoic crystallization ages of around 2180–2041 Ma for orthogneisses of the Mantiqueira Complex and two metamorphic events at 2100 Ma and 560 Ma.

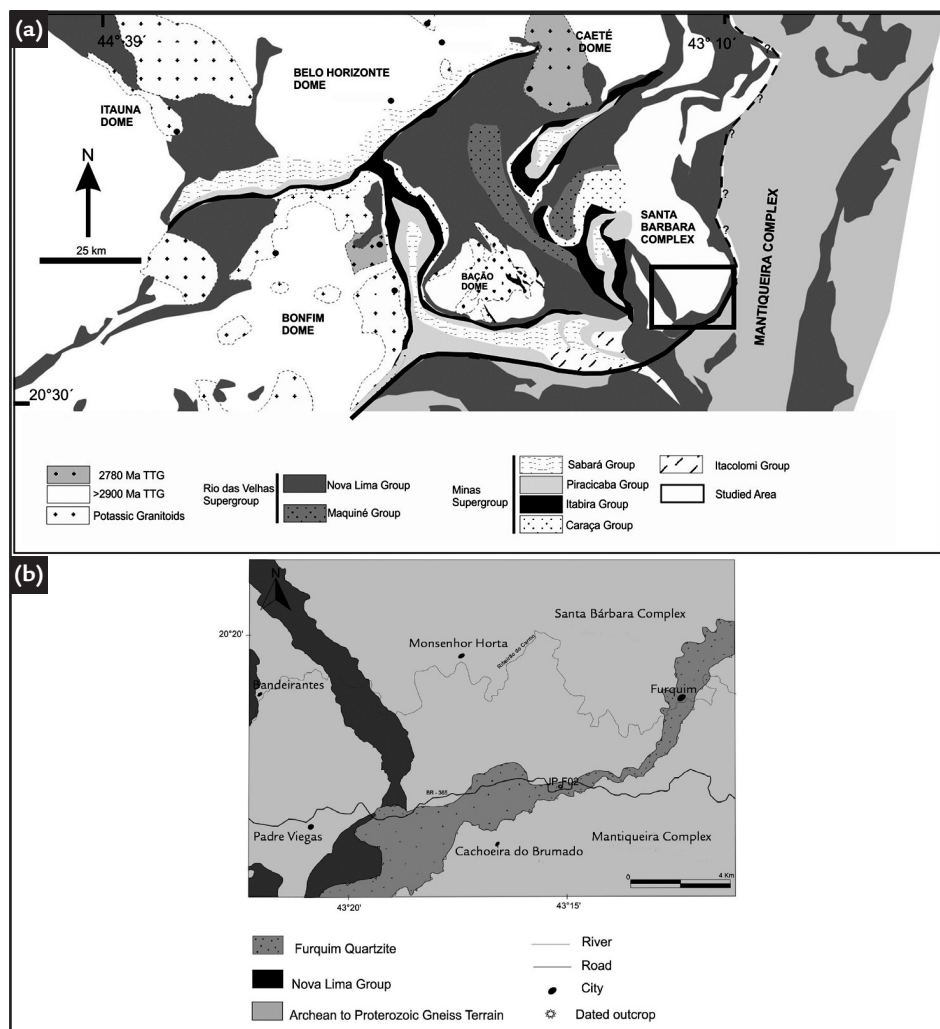


Figure 1
 a) Geological map of the Quadrilátero Ferrífero (modified from Lana *et al.*, 2013). Rectangle: location of the study area;
 b) Geological map of the study area of the Furquim Quartzite (modified from Jordt-Evangelista, 1984) and location of dated sample JP-F02.

Quartzitic unit	Age (Ma)	
	Hartmann <i>et al.</i> , 2006	Machado <i>et al.</i> , 1996
Itacolomi Group	2143±16	2059±58
Sabará Formation	2668±20	2125±4
Moeda Formation (Caraça Group)	2649±11	2651±33; 2606±47
Maquiné Group	-	2877±3
Nova Lima Group	2749±07	2996±38

Table 1
 Youngest zircon grains in the Quadrilátero Ferrífero quartzites corresponding to the maximum ages for the deposition.

3. Material and methods

Thin sections of representative hand samples collected in the vicinity of the town of Furquim and on the highway BR-356 were described on a Leica DM EP microscope at the Departamento de Geologia (DEGEO), Universidade Federal de Ouro Preto.

One sample (sample JP-F02, UTM: 0682390/7746163, sample locality on Fig. 1b) weighing *ca.* 5kg was collected for LA-ICP-MS U-Pb geochronology. Zircons were concentrated making use of a conventional jaw crusher, milling, manual panning and heavy liquids separation. The zircons were hand-picked under a binocular microscope. Approximately 123

zircon crystals were selected and mounted on a 2.5 cm-diameter epoxy mount. The mount was polished and imaged under SEM-cathodoluminescence to accentuate internal growth zoning.

The laser ablation-ICP-MS (LA-ICPMS) analyses were performed using a single collector Agilent 7700 Quadrupole(Q)-ICP-MS and a 213 nm New Wave laser at the isotope/geochemistry laboratory of Department of Geology, Universidade Federal de Ouro Preto. Acquisitions consisted of a 20 s measurement of the gas blank, followed by a 40 s measurement of U, Th and Pb signals during ablation, and a 30 s washout. All

ratios were obtained after averaging the background-subtracted signal (See Romano *et al.*, 2013 and Takenaka *et al.*, 2015, for details on the instrumentation and methodology). Two standards were used during runs: the primary standard GJ-1 zircon (608 ± 1 Ma) and the secondary standard Plesovice zircon (338 ± 1 Ma). The relevant isotopic ratios have been calculated using Glitter data reduction software (van Acherbergh *et al.*, 2001). The U-Pb diagrams were produced using Isoplot 4 software (Ludwig, 2012). The results of the analyses, including data for the primary and secondary standards can be found in Table 2.

4. Results

4.1 Petrography

The main rock type is a strongly folded (Fig. 2a) and sheared muscovite quartzite. Disrupted quartz veins along fold hinges may be confused with pebbles

or cobbles of conglomerates (Fig. 2b). Quartz reaches 70 to 90 vol%, muscovite 5 to 20%, while the accessory minerals hematite, magnetite, kyanite, feldspars,

garnet, zircon, and tourmaline seldom sum 5% (Fig. 2c to 2f, all samples from outcrop in the town of Furquim).

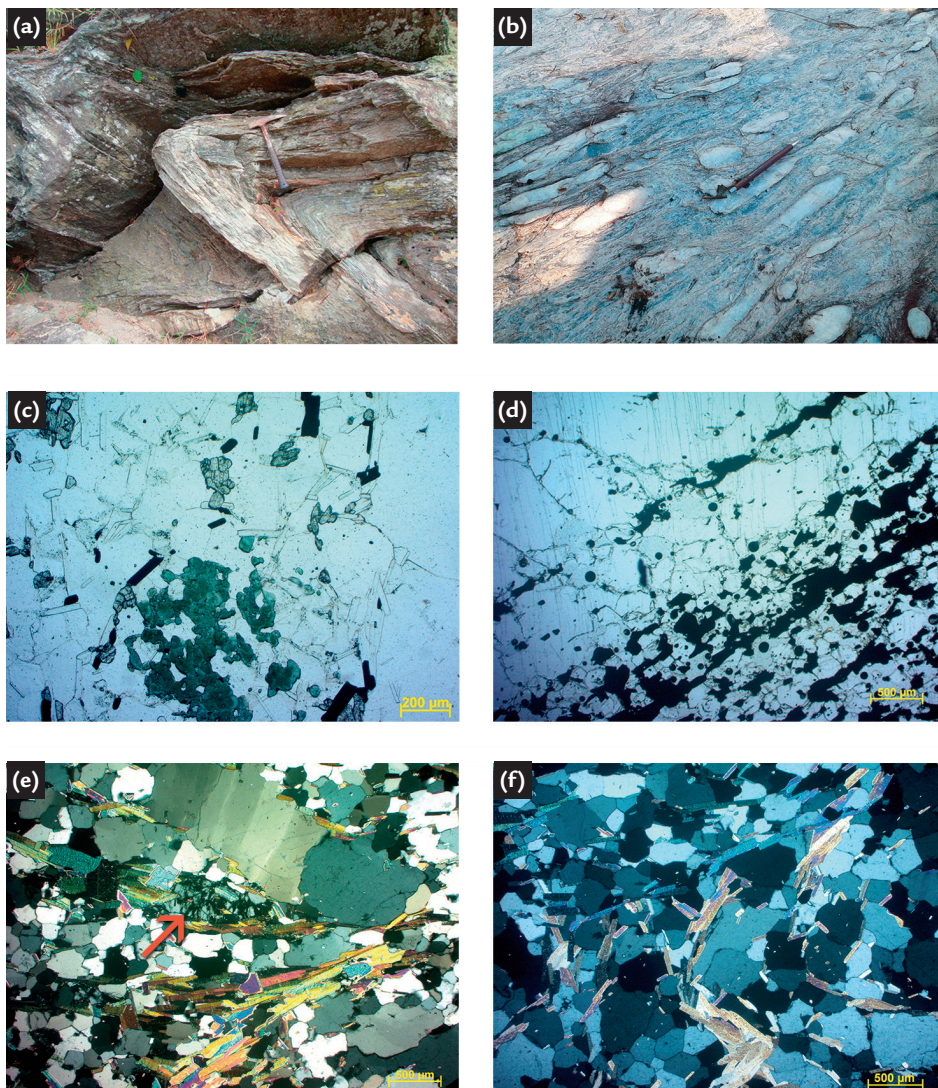


Figure 2
 (a) - Outcrop of the dated folded quartzite (Point JP-F02 on Fig. 1b), view to SSW. See hammer in the center of photo for scale.
 (b) - Disrupted quartz veins, outcrop in the town of Furquim. See pencil in the center of photo for scale.
 (c) - Photomicrograph of muscovite quartzite with tourmaline (greenish), kyanite (gray, strong relief) and hematite (black), N//.
 (d) - Photomicrograph of hematite-rich portion in quartzite, N//.
 (e) - Photomicrograph showing subgrains in deformed quartz vein in muscovite quartzite. Strongly altered feldspar is seen in the center of figure (arrow), NX.
 (f) - Photomicrograph of folded muscovite quartzite, NX. Samples of Fig. 2c to e collected in Furquim. (Photomicrographs from Alvarenga, 2013).

4.2 Geochronology

The extracted zircons from sample JP-F02 measure ca. 100-200 μm , are yellow to brown, slightly rounded and often fractured. Cathodoluminescence images show that most grains present well-defined

oscillatory zoning with some broad zones of intense alteration and radiation damage and no discernible core-rim relationships (Fig. 3). Of the 123 grains mounted on the epoxy disc, 119 LA-ICP-MS analyses

were performed on center and rims of 47 translucent to partly translucent grains. The complete geochronological data set can be found in Table 2.

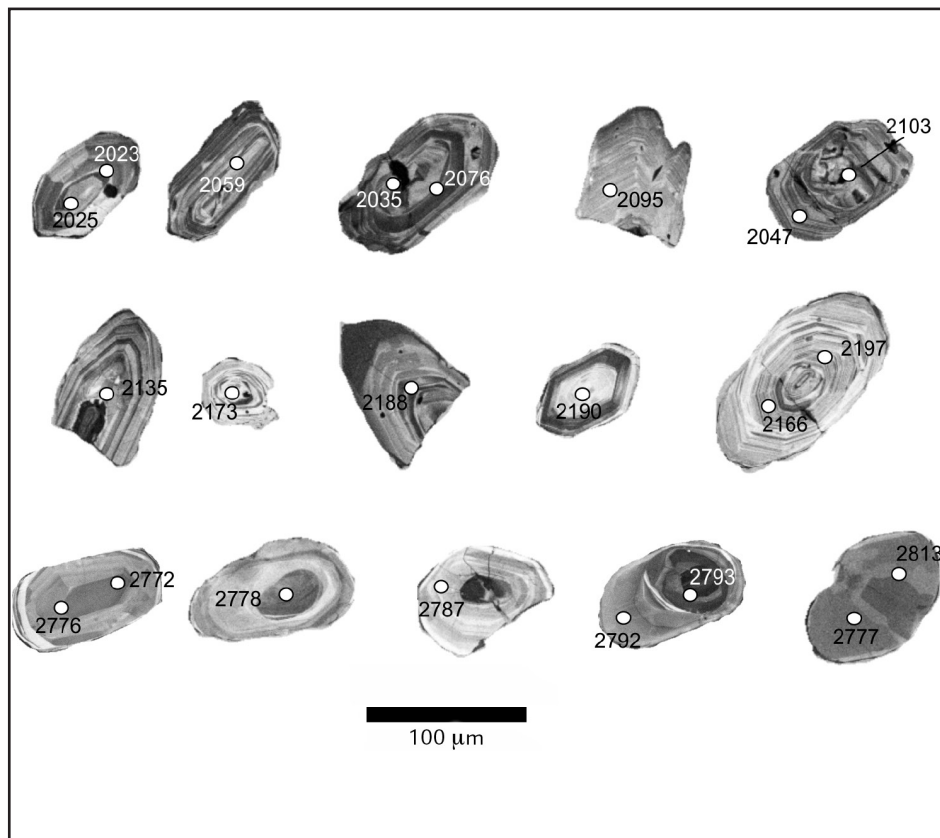


Figure 3
Cathodoluminescence images of analyzed zircons, ages in million years (Ma).

Figure 4 shows concordant to sub-concordant points (63 analyses > 3% concordant) plotted on the frequency histogram and on the concordia. Several

Paleoproterozoic populations ranging between 2.0 and 2.5 Ga (n=31) are distributed along the concordia. The youngest one of 2087 ± 19 Ma old (n=10) defines

the maximum age of sedimentation for the precursor sandstones. An Archean component is represented by ages in the 2.5-3.0 Ga (n=32) range.

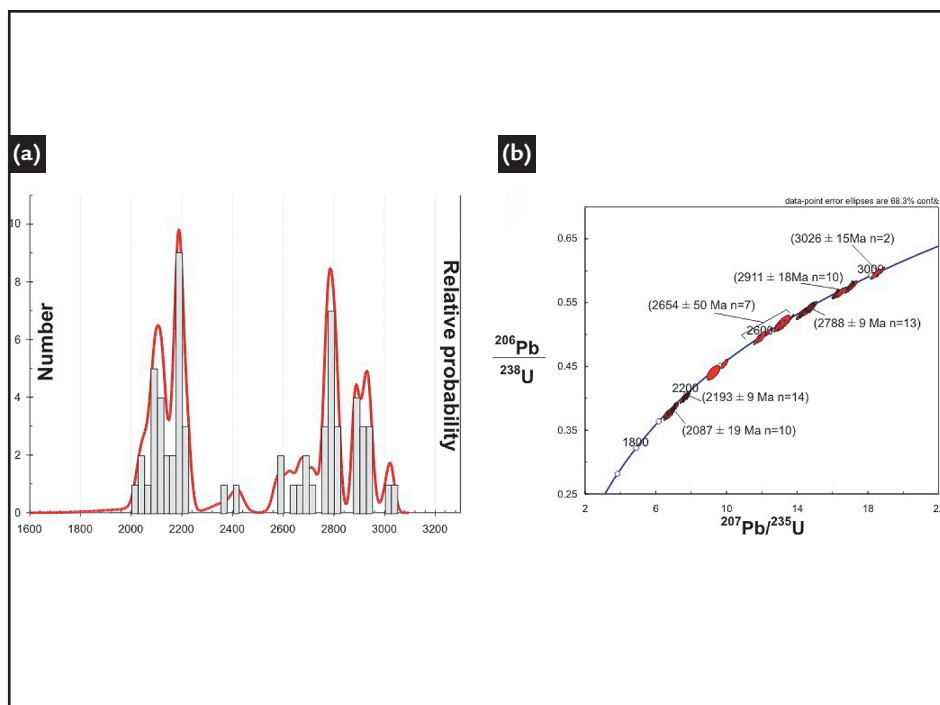


Figure 4
(a) Frequency histogram where two main detrital zircon populations of the Furquim Quartzite can be identified: a Paleoproterozoic population of about 2.2 Ga and an Archean population around 2.8 Ga.
(b) U-Pb Concordia diagram of LA-ICP-MS analyses of zircons (see Fig. 1b for sample locality).

Table 2 - part II
Results of U-Pb LA-ICP-MS analyses for the Furquim Quartzite.

#	Pb204 CPS	Pb207 CPS	U238 CPS	Pb PPM	Th/U	Pb207/Pb206	1s	Pb206/U238	1s	Pb207/U235	1s	Pb207/Pb206	1s	Pb206/U238	1s	Pb207/U235	1s
Padrao Secundário																	
TEST156	6	2851	749890	133,36	0,06	0,05321	0,00061	0,05369	0,00047	0,39393	0,00422	337,6	25,61	337,2	2,85	337,2	3,07
TEST157	1	3190	851437	131,71	0,06	0,05329	0,00054	0,05371	0,00046	0,39472	0,00376	341	22,64	337,3	2,84	337,8	2,73
TEST152	14	2866	755710	128,88	0,06	0,05274	0,00062	0,05373	0,00046	0,39072	0,00426	317,5	26,28	337,4	2,84	334,9	3,11
TEST153	0	2904	757483	131,48	0,06	0,05364	0,00069	0,05374	0,00048	0,39746	0,00481	355,9	28,8	337,4	2,92	339,8	3,49
TEST157	11	2732	701698	129,82	0,06	0,05319	0,00067	0,05374	0,00047	0,39415	0,00464	336,9	28,3	337,5	2,85	337,4	3,38
TEST155	0	2869	759924	134,84	0,06	0,05287	0,00062	0,05378	0,00047	0,39206	0,0043	323,3	26,34	337,7	2,87	335,9	3,14
TEST5	25	2960	786528	127,97	0,06	0,0532	0,00055	0,05389	0,00047	0,39548	0,00385	337,2	23,12	338,4	2,87	338,4	2,8
TEST6	39	3149	854669	127,64	0,06	0,05309	0,0006	0,05392	0,00048	0,39502	0,00423	332,7	25,15	338,5	2,96	338	3,08
TEST8	4	2860	770014	126,45	0,06	0,05258	0,00055	0,05399	0,00047	0,39143	0,00387	310,7	23,67	338,9	2,85	335,4	2,82
TEST36	0	3177	843699	135,6	0,06	0,05324	0,00054	0,05402	0,00047	0,39656	0,00379	338,9	22,8	339,1	2,85	339,2	2,76
TEST158	34	3020	825601	130,26	0,06	0,05244	0,00073	0,05405	0,0005	0,39085	0,00515	304,6	31,19	339,3	3,08	335	3,76
TEST34	11	3173	836267	133,17	0,06	0,05328	0,00053	0,05405	0,00047	0,39722	0,00373	340,5	22,32	339,3	2,86	339,6	2,71
TEST154	25	2977	787230	138,27	0,06	0,05296	0,00058	0,05413	0,00047	0,39534	0,00408	327,2	24,68	339,8	2,88	338,3	2,97
TEST17	6	3018	808971	131,49	0,06	0,05272	0,00054	0,05414	0,00047	0,3937	0,00382	316,8	23,17	339,9	2,87	337,1	2,78
#	Pb204 CPS	Pb207 CPS	U238 CPS	Pb PPM	Th/U	Pb207/Pb206	1s	Pb206/U238	1s	Pb207/U235	1s	Pb207/Pb206	1s	Pb206/U238	1s	Pb207/U235	1s
Padrao Primário																	
TEST90	0	3450	443459	139	0,02	0,0604	0,0006	0,0978	0,0009	0,8145	0,0078	618	22	602	5	605	4
TEST9	9	3826	492675	135	0,02	0,0602	0,0006	0,0981	0,0008	0,8136	0,0076	610	21	603	5	605	4
TEST20	0	3706	473455	140	0,02	0,0606	0,0006	0,0981	0,0008	0,8195	0,0077	624	21	603	5	608	4
TEST39	0	3496	451285	142	0,02	0,0599	0,0006	0,0982	0,0008	0,8110	0,0077	599	22	604	5	603	4
TEST159	0	3026	390152	134	0,02	0,0598	0,0007	0,0984	0,0009	0,8112	0,0084	595	24	605	5	603	5
TEST4	9	3665	477484	135	0,02	0,0593	0,0006	0,0985	0,0009	0,8052	0,0076	577	22	606	5	600	4
TEST121	2	3015	401258	138	0,02	0,0579	0,0006	0,0986	0,0009	0,7868	0,0079	524	24	606	5	589	4
TEST3	1	3670	473661	133	0,02	0,0601	0,0006	0,0986	0,0009	0,8168	0,0076	606	21	606	5	606	4
TEST138	8	3042	391993	135	0,02	0,0597	0,0007	0,0987	0,0009	0,8125	0,0083	593	23	607	5	604	5
TEST31	0	3570	465714	140	0,02	0,0589	0,0006	0,0988	0,0009	0,8029	0,0076	565	22	607	5	599	4
TEST72	0	3376	430147	141	0,02	0,0603	0,0006	0,0988	0,0009	0,8221	0,0079	616	22	607	5	609	4
TEST137	0	3169	399681	135	0,02	0,0609	0,0007	0,0989	0,0009	0,8307	0,0084	636	23	608	5	614	5
TEST150	12	3064	391862	136	0,02	0,0599	0,0007	0,0990	0,0009	0,8183	0,0084	601	23	609	5	607	5
TEST21	13	3735	472705	140	0,02	0,0606	0,0006	0,0990	0,0009	0,8278	0,0077	626	21	609	5	612	4
TEST33	0	3633	466133	140	0,02	0,0597	0,0006	0,0992	0,0009	0,8165	0,0077	593	21	610	5	606	4
TEST38	0	3521	449359	142	0,02	0,0599	0,0006	0,0993	0,0009	0,8208	0,0078	601	22	610	5	609	4
TEST105	10	3131	403413	139	0,02	0,0593	0,0006	0,0994	0,0009	0,8121	0,0081	576	23	611	5	604	5
TEST32	0	3608	460841	140	0,02	0,0598	0,0006	0,0994	0,0009	0,8203	0,0077	598	22	611	5	608	4
TEST122	10	3080	394792	138	0,02	0,0595	0,0006	0,0995	0,0009	0,8168	0,0082	587	23	611	5	606	5
TEST160	11	3043	384716	137	0,02	0,0603	0,0007	0,0995	0,0009	0,8272	0,0085	613	24	612	5	612	5
TEST10	0	3739	478524	137	0,02	0,0596	0,0006	0,0996	0,0009	0,8188	0,0077	590	22	612	5	607	4

5. Discussion

The age pattern of 31 zircons reveals a high concentration of ages in the 2.0-2.5 Ga range, with a mode of around 2.2 Ga and an important Archean component. This pattern is similar to that obtained for the Itacolomi Group by Machado *et al.* (1996), thus suggesting that the Furquim Quartzite detrital sequence belongs to this Group. This interpretation is also supported by the obtained minimum age of 2087±19 Ma that is identical to the minimum ages of 2059±58 Ma obtained by Machado *et al.* (1996) and that of 2143±16 obtained by Hartmann *et al.* (2006) for the Itacolomi samples.

6. Concluding remarks

The results of the geological, petrographic and geochronological investigation of the Furquim Quartzite allowed for a conclusion that the ridge located southeast of the Quadrilátero Ferrífero is possibly an allochthonous unit. The postulated correlation to the Archean Maquiné quartzite as presented in the regional geological map

The large number of zircon grains of Paleoproterozoic age indicates that the main sediment sources for the Furquim Quartzite are terrains generated during the Transamazonian Orogeny. The Archean ages indicate contribution of the gneissic and greenstone terrains.

The youngest Paleoproterozoic population dated at 2087±19 Ma defines the maximum age of deposition of the precursor sandstones of the Furquim Quartzite.

Quartzite ridges belonging to the Itacolomi Group are more widespread than supposed so far. Besides the locus tipicus near the city of Ouro Preto, the

huge mass of quartzite, approximately 1400 m thick, occurring at Serra de Ouro Branco, southern Quadrilátero Ferrífero, previously correlated to the Tamanduá Group or to the Moeda Formation was dated by Machado *et al.* (1996) and correlated to the Itacolomi Group.

The occurrence of a ridge of quartzite belonging to the Itacolomi Group farther east from its locus tipicus near Ouro Preto is probably due to the action of a tectonic event as indicated by its allochthonous nature. Other studies concerning the structural complexity of Furquim Quartzite are necessary to elucidate its geological evolution.

7. Acknowledgments

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published by CPRM (Baltazar *et al.*, 1993) is not supported by the results of U-Pb geochronology on zircon by LA-ICP-MS, which defined several detrital zircon populations ranging from Paleoproterozoic to Archean age. The youngest population dated 2087±19 Ma defines the maximum age for the sedimentation of the precursor sand-

stone. This age is similar to the age of the youngest zircon population of the Itacolomi Group in the Quadrilátero Ferrífero dated by Machado *et al.* (1996) and Hartmann *et al.* (2006). Therefore, it is possible to conclude that the Furquim Quartzite can be stratigraphically correlated to the quartzites of the Itacolomi Group.

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